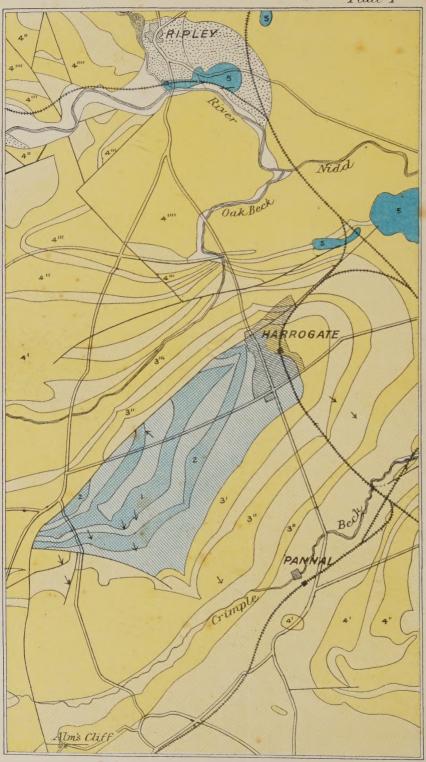




FRED. J. DIXUN
A.M. INST. C.E.
WATERWORKS ENGINEER & MANAGER



# MEMOIRS OF THE GEOLOGICAL SURVEY ENGLAND AND WALES

#### EXPLANATION OF SHEET 62.

(93 NW. OLD SERIES.)

# THE GEOLOGY OF THE COUNTRY NORTH AND EAST OF HARROGATE.

BY

C. FOX-STRANGWAYS, F.G.S.

SECOND EDITION.

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#### PREFACE

Since the first edition of this Memoir was published, thirty-four years ago, additional information as to the geology of the district has been obtained, especially with reference to the glacial deposits of the Vale of York. The physical history of the district as revealed by these deposits has been the subject of important researches by Prof. Carvill Lewis, Prof. Kendall, and others, and advantage has been taken of the necessity of a new edition of the Memoir to refer to these researches.

In the preparation of this edition we have fortunately been able to secure the services of the original author, who has also added new matter relating to the structure of the area, the alteration of the river courses, the history and origin of the celebrated Harrogate Springs, and an extensive bibliography dealing with the geology of the district, and all subjects relating to the springs, except such as are of a purely medical character.

Our thanks are due to Mr. Godfrey Bingley for the photographic illustrations of some of the more remarkable features of the district.

J. J. H. TEALL,

Director.

Geological Survey Office, 28, Jermyn Street, London. 2nd September, 1908.

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#### THE GEOLOGY

#### OF THE COUNTRY NORTH AND EAST OF

### HARROGATE.

#### CHAPTER I.

#### INTRODUCTION.

The new series Sheet 62, known previously as quarter-sheet 93 N.W. (old series), may be said to comprise the valley of the Ouse from Ripon to York, or, roughly speaking, the central part of the Vale of York. The area contained in the map is 216 square miles. The chief towns are Harrogate, Knaresborough, Boroughbridge, and Easingwold, with the southern suburb of Ripon; the map also extends, on the south-east, to within three miles of the

city of York.

The principal rivers are the Ouse, the Ure, the Swale, and the Nidd, which all rise in the high ground in the north-west of the county, and uniting into one stream, in this map, enter the German Ocean by the estuary of the Humber. Thus the river system belongs entirely to the basin of the Ouse, the name by which the united streams are known below their junction with the small rivulet flowing from Great Ouseburn. The direction of all the streams is more or less easterly, and is due to the general inclination of the Permian and Triassic rocks being in that direction. The earth movements which have thrown the older rocks into anticlinal and synclinal folds have had no influence in determining the lines of drainage. It will be seen later on that this easterly drainage was much interfered with during Glacial times, and that the original courses of many of the streams were changed during that epoch.

The ground rises gradually on either side of the great central valley to the north-east and west, but principally towards the west, on which side the highest ground is at Harlow Hill near Harrogate, and How Hill, near Fountains Abbey, both of which are about 600 feet above the sea. The general dip of the several formations is to the north-east, although in the western part of the district the Palæozoic rocks are thrown into a series of anticlinal and synclinal folds striking north-east and south-west,

the most remarkable of which is that at Harrogate.

All the physical features of the country may be shown to be a result of its geological structure. Thus, the hard sandstones and soft shales of the Carboniferous rocks form the high ground and deep valleys in the western part of the map; these are succeeded by a lower, but well-marked ridge of Magnesian Limestone rising to an average height of between 200 and 300 feet above the sea; beyond these again are the soft marls and sandstones of Permian and Triassic age, forming part of the great central valley of Yorkshire.

A small area at the north-east corner of the map, where the ground begins to rise again in that direction, is occupied by the shaly sandstones of the Middle Lias, the soft marks and shales of the Keuper and Lower Lias being entirely hidden by an immense

thickness of alluvial deposits.

The country under notice is almost entirely agricultural, there being no mines or manufactures of any extent. The principal and most prosperous place in the map is Harrogate, the importance of which is entirely due to its geological position. The sharp uplift of the lowest rocks that occur in the district, which takes place at this point, has given rise to such a variety of mineral waters that during late years Harrogate has become one of the best known health resorts in the country. The intimate connection between the geological structure of the district and the origin of these springs is of so much interest that a special chapter has been devoted to this subject.

#### CLASSIFICATION AND TABLE OF STRATA.

Since the publication of this map in 1874 it has been shown by Dr. Wheelton Hind that the important series of strata between the Millstone Grit and the Carboniferous Limestone of Skipton and beyond are not the homotaxial equivalent of the Yoredale Shales of north-west Yorkshire. Consequently the term Yoredale has been withdrawn for the description of these beds in the present memoir, as it is probable that they should be referred to the southern series rather than to those of Wensleydale.

The red grits at the base of the Magnesian Limestone were at one time considered to be of Permian age, but the error was pointed out long ago, and has been fully referred to elsewhere, so that it is not necessary to go into the subject further.

The age and separation of the sandstones forming the greater portion of the Trias of this part of the country has always been a matter of some doubt, and no further light has been thrown on

the subject of late years.

The origin of the Glacial beds and the physical geography of that period has been rendered much clearer by the work of Professor P. F. Kendall in this county since the first edition of the memoir was written.

The following list gives in descending order all the divisions and subdivisions of the strata that occur within the area:—

#### Table of Strata.

RECENT AND POST- Alluvium and River Terraces.  GLACIAL - Warp and Lacustrine Clay, Sand and Gravel.  GLACIAL - Glacial Clays, Sands and Gravels.					
JURASSIC	Upper Lias. Middle Lias. Lower Lias.				
TRIAS -	Rhætic. Keuper. Bunter.				
PERMIAN	Upper Magnesian Limestone. Middle Marl. Lower Magnesian Limestone. Lower Marl and Sands.				
Carboniferous -		Plumpton Grit. Shales. Cayton Gill Beds. Shales. Follifoot Grit. Shales. Kinder Scout Grit in three beds.			
	Harrogate Roadstone Series. "Yoredale Rocks" of previous edition of memoir.	Shales. Roadstone. Shales. Grit. Shales.			

All the subdivisions in the Table are represented by distinct colours except the Drift, which at the time this country was surveyed was not shown on the map. A considerable amount of Boulder clay and gravels, however, occurs in the district, and obscures the rocks that lie beneath, so that many of the boundaries are very uncertain. For this reason many of the boundaries between the Permian Marls and Limestone are very doubtful; while that between the Permian and Trias is entirely conjectural, the junction between these formations not having been observed at any point, although by aid of well sections its approximate position is not difficult to fix. The boundary between the Trias and the Jurassic rocks is also entirely concealed by superficial beds, but its approximate position and the general structure of these underlying rocks can be made out from a study of the country further to the east, which shows that the strata are much disturbed by lines of fault These lines of fault are not shown beneath the superficial beds on this map, but their general position may be gathered from the four miles to an inch map of this district.

#### CHAPTER II.

#### CARBONIFEROUS.

#### HARROGATE ROADSTONE SERIES.

The lower Carboniferous rocks of Yorkshire were formerly divided by Professor Phillips into the two main divisions of Lower Scar Limestone and Yoredale Rocks; but it has since been pointed out by Dr. Wheelton Hind that the second of these names cannot be applied to the series in the southern part of the county, from the fact that the beds which occur in the typical district of Yoredale or Wensleydale are on a different horizon to those of similar character in the south. Dr. Hind has shown that the true Yoredale beds are the homotaxial equivalent of the upper part of the Carboniferous Limestone, and consequently that it is erroneous to apply this name to the beds above the Limestone in the southern area. For these latter he proposes the name Pendleside series, from Pendle Hill where the beds are thickest and best exposed.

The district we are at present concerned with lies midway between these two places, and it is doubtful to which type the beds at Harrogate should be assigned. The typical Pendleside series dies out at about this latitude, and as the rocks at Harrogate contain few organic remains except Crinoids, there are no palaeontological data by which the horizon of these beds

can be definitely determined.\*

The nearest exposure of these rocks is in the valley of the Wharfe about Bolton Abbey; but the beds, which here crop out beneath Beamsley Beacon, are so different in character from those of the Harrogate inlier that they throw no light on this section.

At Harrogate these beds consist, for the most part, of shales with a few bands of cherty limestone, and one tolerably massive bed of sandstone near the base. This sandstone is well seen in a line of quarries on How Hill near Harrogate; it is here a thin-bedded but somewhat hard sandstone, of a whitish colour, and contains on its surface a white efflorescence. On the northern side of Harlow Hill this lock is more decomposed, and near the surface has a rubbly and somewhat sandy or gritty nature.

The limestone is an exceedingly hard silicious rock containing Encrinites in great numbers, in fact in many cases it is rather an aggregated mass of the silicified remains of these than a limestone proper; when by exposure to air and moisture the calcareous portion of the rock becomes dissolved out, it has very much the appearance and texture of pumice stone, and especially is this the case along the angles of flexure when the beds are much contorted.

<sup>\*</sup> Professor Phillips considered the Harrogate roadstone to be the equivalent of the main or 12-fathom limestone at the top of the Yoredale series. Quart. Journ. Geol. Soc., vol. xxi., p. 234.

Where this group of strata crops out to the west, there is no rock exactly corresponding to the Harrogate Roadstone, and it was probably formed under somewhat local and peculiar circumstances.\*

The best sections of this rock are in quarries at Shaw Green, Beckwith House, and behind the Cold Bath Road, Low Harrogate, at all of which places it forms an excellent roadstone and has

been quarried for that purpose.

In this district these rocks are brought up along an anticline running in a N.E. and S.W. direction, and it is owing to this peculiar structure that we have their outcrop in the country under consideration; whenever this cherty limestone is exposed to view it is either contorted or highly inclined having evidently been subjected to great disturbance.

#### MILLSTONE GRIT.

The lower portion only of this group of strata crops out in the district under consideration; the beds which appear in the map being the equivalents of the Third Grit, and Fourth or Kinder

Scout Grit, of Derbyshire.

The Kinder Scout Grit consists of three separate beds of sandstone which, with the intermediate shales, have a thickness of nearly 1,400 feet. Both the lowest and highest of these beds are coarse, massive gritstones, frequently containing pebbles of quartz, and capable of being split by means of joints into blocks of a considerable size. The upper bed has been largely quarried for building-stone at Hookstone Wood, and several places to the south of Harrogate, and also at Birk Crag to the west and near Walker Road to the north of the town.

The middle bed of sandstone belonging to this group is not nearly so clearly defined as the other two. It consists of a thin flaggy sandstone which becomes more massive towards the base; it forms no good feature anywhere throughout the district, and not being well exposed has, apparently, been overlooked by

previous observers.

That the strata which comprise this group have a considerable thickness is proved by the fact that a borehole, which was put down at the southern end of Bogs Lane, near Starbeck, near the outcrop of the upper bed failed to reach the sandstone next below, although sunk to a depth of 437 feet. This section, which is given in the Appendix, page 54, was sunk almost entirely in shale, while others nearer Starbeck were in the upper part of this rock.

The Third Grit of this area consists of many separate beds of sandstone; but several of these, owing to local thinning out, are difficult to follow for any distance. They are however divisible into three distinct series which can be easily recognised throughout the district. These which are named after the localities where they are best exemplified are the Follifoot Grits, the Cayton Gill Beds, and the Plumpton Grits.

The Follifoot Grits, so named on account of their being well seen along the edge of the hill south of the Crimple Valley, known as Follifoot Moor, usually consist of two beds, which with

<sup>\*</sup> See also W. H. Hudleston, Proc. Geol. Assoc., vol. vii., p. 426.

the intermediate shale have a thickness of about 200 feet; they are hard compact sandstones which in small fragments have very much the appearance of loaf sugar. To the north of Harrogate these beds form the remarkable ridge of Long Crag north of Oak Beck, and are again seen at the Gas Works, and at several places near Harrogate End. They also outcrop in the bed of the river Nidd below Bilton, but have not been separately mapped, as their course is too obscure to allow it to be done with accuracy.

At Bilton coal was formerly worked in these grits and was used for burning the Magnesian Limestone close at hand, but it was never of much importance. It occurs in two seams separated by about three or four yards of shale. The uppermost of these, which was the one principally worked, had a thickness of 3 ft. 2 in.; the lower bed, which in some places was as much as 2 ft. 9 in. in thickness, appears to thin away altogether on the northern side of the working. The coal has been worked out to the east, but on the west or dip-side it was discontinued on account of water.

In the little stream, west of Thornton Moor House, this coal was again met with at a depth of 60 ft., where it had a thickness of 18 in., but only 4 in. of this was good. It has also been worked

further west at several places in the next map.

The Cayton Gill Beds, so named as the best sections of them are exposed at two or three places in that valley, are not of any great thickness, but are remarkable from their being highly fossiliferous, an unusual occurrence in beds of Millstone Grit age. In Cayton Gill these beds consist of the following sub-divisions:—

Flags with Encrinite remains. | Hard bed.

Beds full of Brachiopoda. Black shale with fossils.

The hard beds have been quarried as a road-material at Rhodes House near Follifoot, Four Lane Ends and Saltergate Hill northwest of Harrogate, at several places on and near Scarah Moor, and by the road leading from Fountains Abbey to Sawley.

The following fossils were collected by the Survey from these

beds at the several localities named:-

	Four Lane Ends, Harrogate.	Hamps-	Railway Cutting Hamps- thwaite.	Fountains Abbey.
	1			
Schizophoria resupinata				1
(Mart.)	×			×
Spirifer bisulcatus J. de C.				
Spirifer striatus? (Mart.)	×	×		T-No-based
Orthototog evenistric (DL'77)	×			
Orthotetes crenistria (Phill.)			×	
Productus sp			×	×
Rhynchonellid	-		×	
Grammatodon cf. cancellatus				
(Mart.)	×		_	
Stroboceras sulcatum (J. de	1	1		1
C. Sow.)	×	×		
Temnocheilus tuberculatus	1			
(J. Sow.)		×	_	

A rather fuller list is given by the Rev. J. S. Tute, *Proc. Yorkshire Geol. Soc.*, vol. ix., p. 266, who also records an *Orthoceras* and a *Lingula* from the Millstone Grit of this district, *Ibid*, vol. vii., p. 82, and vol. ix., p. 425. *Pleuronautilus nodoso-carinatus* (Röm.) was found by Mr. Dixon in the shales that were excavated in making the Scargill reservoir, which are about at this horizon. The specimen is now in the British Museum.

From these beds in this neighbourhood also Dr. Wheelton Hind has lately recorded the following much fuller list.\*

		Clint Quarries.	Near Pateley Bridge.	Hazel Hill, Sawley.
Fenestella sp	_		×	
Chonetes cf. laguessiana -	-	×		×
Derbya sp	-	×	×	×
Productus cora [late mutation	n] -	×	×	×
,, longispinus	-	×	×	×
", scabriculus	-	×	P-1-1-1	_
Rhipidomella michelini -	-	TRANS. IS	×	
Schizophoria resupinata -	-	×	×	×
Seminula ambigua	~	×	×	
Spirifer bisulcatus	- ]	×	- Andreada	
Spiriferina cristata	-	×	×	
Amusium concentricum -	-		×	
Aviculopecten dissimilis -	- '	×		×
,, semicostatus-	-			×
" stellaris -		-	_	×
sp	-	P-100	×	
Cypricardella sp		_	×	
Edmondia maccoyi	-	_		×
" rudis	-	_	_	×
r. " sp	-		×	
Leiopteria laminosa	-	×		
,, squamosa	-		×	
sp	- }		_	×
Lithodomus jenkinsoni	-	_	-	×
Mytilomorpha sp	-	_		×
Palæolima sp	-	~-		×
Parallelodon obtusus	-	×	×	
Protoschizodus curtus -	-			×
Pterinopecten whitei -				×
Sanguinolites sp.		×		×
Tellinomorpha cuneiformis -		^		×
Funhamus en	_		×	×
Euphemus sp	-	_	×	_
Macrochilina sp	-			×
Macrochilina sp Naticopsis sp	_			×
Ephippioceras bilobatum -		_		×
Orthoceras sp	_	×	_	×
Stroboceras sulcatum -	_	×	×	×
Petalodus acuminatus-	-	×	_	
Rhizodopsis sauroides -	-	×	_	_
Tentzouopata aautotuoa		,		

<sup>\*</sup> Naturalist 1907, pp. 92, 93.

The Plumpton Grit, the highest of the Millstone Grit measures occuring in this district, consists of two beds having a total thickness of 150 feet, and composed of thick-bedded, coarse, and exceedingly massive grits. The base of the rock is often a true conglomerate containing rolled pebbles of quartz, and is very ferruginous. This rock was formerly considered as belonging to the Lower New Red sandstone of the Permian series, and to be the equivalent of the Rothliegende of the Thuringerwald in Germany,\* but the sections showing the junction between the two formations are so clear in this neighbourhood that it is difficult to understand how these grits could ever have been supposed to be other than of Carboniferous age.†

The junction of the grit with the Magnesian Limestone above is well seen at several places in the southern part of the map; but, undoubtedly, the finest sections are those exposed along the banks of the Nidd below Knaresborough. Here, although the sections are not always quite clear at first sight, a decided unconformity between the two series of strata can be traced; the denuded surface of the coarse red grit of Plumpton, which is exposed in a series of fine sections, being covered directly by the Magnesian Limestone. This is more particularly the case below what is called Abbey Crag, where there is a prominent boss of red grit which rises several feet into the surrounding limestone.

At Newsome Bridge and St. Helen's quarries, in the extreme south of the map, the junction is also well shown; at the latter of these there is a small patch of red marl, with a little greenish layer below the limestone, which may be the equivalent of the lower marls.

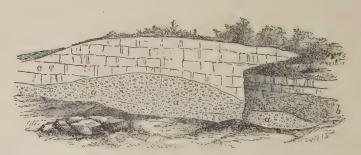


Fig. 1. Newsome Bridge Quarry. Drawn by J. C. Ward.

- a. Whitish grit (Millstone Grit). About x it contains, in quite its upper part, little horizontal patches of a soft unctuous marl.
  - b. Grit and limestone confusedly mixed together.
  - c. Yellow limestone, with included fragments of grit at its base (Permian).

<sup>\*</sup> Trans. Geol. Soc., ser. 2, vol. iii. p. 37; Quart. Journ. Geol. Soc., vol. xxi., p. 234; Murchison's Siluria, 3rd Ed., p. 349.

<sup>†</sup> Geol. Mag., vol. iii., p. 49; Proc. Yorkshire Geol. Soc. for 1867, vol. iv., p. 568; Quart. Journ. Geol. Soc., vol. xxv., p. 291.



Photographed by Godfrey Bingley, Headingley, Leeds.

BANKS OF THE RIVER NIDD AT KNARESBOROUGH. MAGNESIAN LIMESTONE RESTING ON MILLSTONE GRIT.







Photographed by Godfrey Bingley, Headingley, Leeds.

THE DEVIL'S ARROWS, BOROUGHBRIDGE.

MONOLITHS OF MILLSTONE GRIT.



Fig. 2. St. Helen's Quarry. Drawn by J. C. Ward.

a. Coarse purplish grit (Millstone Grit).
b. Red marls, 5 ft., with a little greenish layer at top, below which is occasionally a thin layer of fine grit (redeposited) (Permian).

c. Soft yellow limestone (Permian).

The Plumpton Grit, on account of its great durability as well as the comparative facility with which it may be worked, is largely used as a building-stone, and is quarried at a great number of places along its out-crop. South of the Nidd the best sections are those at Plumpton, near Knox Mill and near Killinghall Mills; north of the Nidd it is well seen at Knaresborough, Scriven, Lingerfield, Ripley, South Stainley, Dole Bank, and Fountains.

It is from this grit that the remarkable blocks known as the Devil's Arrows have come. These huge monoliths consist of three masses of gritstone, the origin and purpose of which has long been a puzzle to antiquaries. They are situated in some fields near the lane leading from Boroughbridge to Roecliffe, about a quarter of a mile from the former place. The blocks stand north and south and are deeply furrowed by weathering; the loftiest one is about 22 feet above the ground. There were formerly four of these stones standing in the 16th century, but one fell and was carried away to form the foundation of a bridge over the Tutt Beck. They are about eight miles in a direct line from Plumpton, and about six from South Stainley and Lingerfield, the nearest places where such a grit exists in situ.\*

These grits, probably from their unequal hardness and thickbedded nature, have a tendency to weather into very peculiar and fantastic shapes, as may be seen more especially at Brimham and Plumpton, as well as either side of the road between the latter place and Spofforth, where they stand up in the centre of the fields in the most picturesque manner; they are usually of a purple colour, which may, possibly, be due to the peroxidation of the iron they contain, by the action of carbonated water filtering from the limestone above through the porous grit.+

† On the colouring of these grits see papers by J. C. Ward, Quart. Journ. Geol. Soc., vol. XXV., p. 295, and Geol. Mag., vol. IX., p. 389, and J. Lucas, Geol. Mag., vol. IX., p. 338.

<sup>\*</sup> These remarkable monoliths are noticed in nearly all the Histories and Guide Books referring to the district, but more especially by W. Camden, 1586, Gibson's edition, 1720, fol. 716; T. Gale, 1709; J. Lleland, 1710; W. Stukeley, 1724; F. Drake, 1736; E. Hargrove, 1769; T. Pennant, 1804; and J. Pbillips, 1853. The full references to these are given in the Bibliography. There are also three drawings of these stones in the print room at the British Museum, Kaye v., 15-17. See also footnote, page 93.

The Millstone Grit at several places is covered with a coating of calcspar, and sometimes small pieces of the rock are cemented together by the same substance; as at Fountains Abbey and

other places.

There are some large detached blocks of this grit at Haddockstones, which have been supposed to be erratics brought down by a glacier from Hebden Wood and Sawley; but there is little doubt that this rock is here *in situ*, or if it has been moved at all it is not more than can be accounted for by slipping.\*

<sup>\*</sup> These stones have been so named from their resemblance to shocks of corn locally termed "haddocks."



PULIFFON HOUSE. SHOWING THE CHRESTS WHATBERING HE THE PURIFIC OF THE MILLSTONE GRIT, Photographed by Godfrey Bingley, Headingley, Leeds. (Copyright.)



#### CHAPTER III.

11

#### PERMIAN.

The Permian rocks of this district are capable of division into four separate alternations of strata, which are to a certain extent unconformable to each other, although this is not always very clearly seen in the country under consideration.

They consist of the following alternations of marl and lime-

stone in descending order.

Upper Magnesian Limestone. Middle Marl. Lower Magnesian Limestone. Lower Marl and Sands.

In the country to the south an upper bed of marl was exposed in a few places, but it is extremely doubtful if such a bed occurs in this district, or could be separated from the overlying Trias.

These beds occupy, as a whole, a broad strip of country in the western half of the map, running from the banks of the Nidd south of Knaresborough, to the neighbourhood of Studley Park, to the west of Ripon.

Lower Marl and Sands:—In the St. Helen's quarry there are five feet of red marl occupying a depression in the underlying grit which probably indicates the existence of these marks in the district, but the outcrop, which must be very limited, is not

capable of being traced.

At Brearton some beds of marly sandstone were seen, which are also probably of this age. They are composed of soft brickred sandstones, excessively like the Bunter sandstone, but unlike rocks of that age they are interstratified with thin seams of white unctuous clay, resembling fuller's earth. The subjoined rough sketch was taken from the sandpit at the back of the village, and may give some idea of the nature of this deposit:-

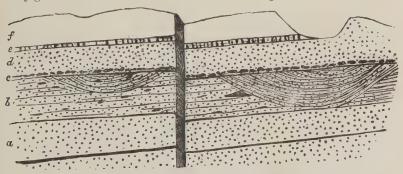


Fig. 3. Sand-pit near Brearton.

a. Red sandstone. b. Red sandstone, much false-bedded, earth. with fragments of red marl and d. Red sandstone. white clay.

c. Band of fragments, chiefly fuller's

e. White unctuous clay.

f. Red sand.

The beds dip to the north at an angle of about 10 degrees; they are much broken up and false-bedded, and appear to have been much disturbed during their formation, having some-

what the character of a shore deposit.

The true geological age of this rock may be a matter of dispute, but we prefer to class it with the Permian on account of its peculiar lithological character, although we believe a similar deposit has not been observed anywhere throughout the range of the Permian measures in this part of England.

At the base of the Magnesian Limestone there occasionally occur certain sands which appear to be formed by the decomposition of the limestone itself; they were first noticed by the late Professor Sedgwick in his paper on the Permian rocks of

this part of England.

These sands were seen in a pit on the western side of the out lying patch of limestone at New Covert, near Scriven, where the following section was shown:-

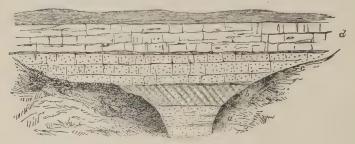


Fig. 4. Sand-pit at New Covert, near Scriven.

a. Fine white sand, with thin yellow clayey bands.
b. Sand with hard veins, which form ridges inclined to the plane of bedding.
c. Yellow sandy limestone.
d. Coarse and rubbly limestone.

a and b are of a slight pink colour towards the top.

The bottom of the sand is not reached, so we could not arrive at its thickness, but it cannot be very great, as coarse conglomeratic beds of red grit are seen a short distance to the south, dipping in the direction of the limestone. These sands have been used at Glass Houghton, Pontefract, Harthill, and other places for moulding purposes.

Lower Limestone.—This is the principal rock of the Permian measures, having an average thickness of about 175 feet. It is usually a soft, thick-bedded, yellowish, sandy limestore, but its character varies somewhat in different localities; frequently it is very concretionary, and contains small cavities lined with bitter spar, causing the surface of the rock to weather into curious shapes.

The rock is well exposed in various quarries, but the best sections are to be obtained from the banks of the Nidd, near





Photographed by Godfrey Bingley, Headingley, Leeds. (Copyright.)

GRIMBALD CRAG, KNARESBOROUGH. MAGNESIAN LIMESTONE ON MILLSTONE GRIT.

Knaresborough, in the railway cutting at Wormald Green, and the banks of the Skell in Studley Park. Some of these sections show that the rock is occasionally much contorted, but this is probably due to local settlement rather than actual disturbance.

The base of the rock, when exposed, is seen to contain fragments of grit and quartz, derived from the underlying Millstone Grit, and has somewhat the character of a conglomerate, being much coarser in texture and harder in substance than the main mass of limestone. This character was well seen in a quarry near Grimbald Crag, at Goldsborough Mill, and other places.

At Markington this conglomerate, which contains waterworn pebbles of limestone and sandstone, is about two or three feet in

thickness.

Nodules of impure silicate of lime, which may have been formed round some organic substance, occur in the limestone at Wormald Green.

At Quarry Moor, near Ripon, the rock is traversed by a series of small faults, which give it a peculiar veined structure, the limestone (probably by pressure) being rendered more compact and of a blue colour for some little distance on either side of the lines of fracture.

The upper part of the limestone is frequently much softer and of a more earthy nature, and in the neighbourhood of Farnham and Burton Leonard has been dug in what are called marl-pits, and spread over the land.

The Rev. J. S. Tute has divided this limestone into four

horizons in the following order\*:—

4. Main mass of limestone.

50 feet of limestone with irregular lumps.
 Yellowish limestone with Axinus obscurus.

1. Lowest beds with Productus horridus.

Mr. Tute considers that the lowest beds of Fountains and Knaresborough are not the lowest beds of other places; but whether the earlier beds were ever deposited at these places or have since been denuded is not stated. It, however, shows that considerable irregularity exists at the base of the Permian measures.

Small caves frequently occur in the limestone and are usually filled with clay and stones which have been washed in from above; this is the case with Mother Shipton's Cave at Knaresborough, and also in quarries near Markenfield Hall and in Studley Park Sometimes the caves form underground passages in which the present streams flow. A good instance is seen in Studley Park, where during the summer the River Skell is entirely swallowed up, and emerges again in the form of a copious spring at a place called Hell Wath, a mile and a half lower down its

These caverns or fissures are probably the cause of the "blowing wells" that occur along the outcrop of the Bunter

<sup>\*</sup> Proc. Yorksh. Geol. Soc., vol. viii., p. 218.

Sandstone, and which have been described at some length by Mr. T. Fairley.\* Strong currents of air flow from the shafts of these wells during a falling period of the barometer, and inward currents are found when the barometer is rising. phenomenon has been observed at Ornhams Hall to the south of Boroughbridge, at Solberge to the south of Northallerton, and at Langton to the north of that town, but it doubtless occurs at other places in the district.

The Lower Limestone when exposed in this area is not very fossiliferous, but in a quarry at Aldfield, close to the north-west corner of the map, and also in one at Well rather further north, the following list of fossils was obtained by the Rev. J. S.

Tutet -

Cythere geinitziana Jones Acanthocladia anceps Schloth. Fenestella retiformis Schloth. Productus horridus Sow.

latirostratus Howse Terebratula elongata Schloth. Streptorhynchus pelargonatus Schloth.

Strophalosia lamellosa Geinitz Spirifera alata Schloth. schlotheimii Camarophoria

von Buch

Schizodus truncatus King rotundatus Brown Bakewellia (Gervillia) antiqua Münst.

ceratophaga Schloth. Monotis speluncaria Schloth. Solemya sp.

Pleurophorus costatus Brown Leda speluncaria Geinitz Turbonilla phillipsii Howse Turbo helicinus Schloth. Nautilus frieslebeni Geinitz

The Middle Marl.—The Middle Marl does not appear to be so thick nor to occupy so large an area as in the map to the south. It consists of red marl and soft red sandstone, which in some sections is excessively like that of Triassic age. There are very few good sections in this rock, the best being those on the banks of the Nidd below Goldsborough Mill and to the west of Bilton Hall; it was formerly visible in the railway-cutting and in the clay pit west of Knaresborough, but these latter are now overgrown and obscured; beyond these places its presence is denoted by the red and clayey nature of the soil. To the north of Knaresborough the country is so thickly covered with Glacial and Post Glacial deposits that there is no evidence of the presence of the Middle Marls, except perhaps near Monkton Mains, where there was formerly an old marl-pit, and where the ground is of a somewhat reddish nature.

These marls in some places are stated to contain large quantities of gypsum, which has been used for making plaster of Paris; but this does not appear to be the case in this district, and the beds of gypsum which occur near here are in higher strata.

The Upper Limestone.—This is, also, not nearly so important a rock as it has been further south, the small patches in the neigh-

<sup>\*</sup> Proc. Yorksh. Geol. Soc. vol. vii., pt. iv., p. 409.; also A. C. G. Cameron. Geol. Mag. 1880, p. 95. + Proc. Yorksh. Geol. Soc. vol. iv., p. 557, and MS. notes.

bourhood of Goldsborough and Knaresborough being all that occur in the map. At these places it is a thin fissile limestone, easily crumbling away on exposure to the atmosphere, and of scarcely any commercial importance. It is well seen in the railway-cutting west of Knaresborough, resting on red marl, and in quarries at Goldsborough where it is fossiliferous.

The following species from Goldsborough have been determined

by Dr. Ivor Thomas:

Mytilus septiter King Schizodus rotundatus (Brown)

The Upper Marl has not been observed in this district.

#### CHAPTER IV.

#### TRIAS.

Owing to the thick covering of Drift deposits which extends across the centre of the district very little is known as to the general relations of the Trias to the beds that lie below. Practically the Permian and Trias form an unbroken series, and although the former is linked paleontologically with the older rocks, lithologically it is more allied to those above. For this reason it is difficult to separate the marks of one formation from the sandstones of the other. In this region, however, the ground is, as we have said, so obscure that this difficulty is not of much account, as the division between the two can be but merely approximate. Whether the Trias is conformable to the Permian or whether the former overlaps the latter there is no evidence to show.

The New Red Sandstone consists probably of the two divisions of Bunter and Keuper Sandstone and Marls, although owing to the thick covering of superficial deposits it is only the former that is actually seen at the surface, and even that at only comparatively few places.

#### BUNTER SANDSTONE.

This is a soft, brick-red, thick-bedded sandstone, containing occasionally thin marly partings, and in one locality several veins of opaque white quartz. The best sections of this rock are to be found in the railway-cutting half a mile east of Cattal Station, and at Aldborough; besides which it is exposed near the bank of the river at Hunsingore, in the bank of the river at Broad Oak Farm, at the Keeper's Lodge north of Hunsingore, at Cattal Station, in the road and lane east of Green Hammerton,\* at Thorpe Green, at Little Ouseburn, and ma field south of Marton. It is also well seen at several places to the north-east of Ripon just beyond the northern boundary of the map. At Aldborough and Ripon it was formerly used as a building-stone, but these quarries are not now in use.

A curious circumstance connected with this rock is its liability to sink and form depressions and hollows in the surface, from 50 to 100 feet across, and occasionally perpendicular shafts of large size. Several of these pits occur in the northern part of the map near the town of Ripon. At Bishop Monkton one of these sinkings occurred about 40 or 50 years ago, which carried away a stack that was being made at the time; but the most remarkable one which we have seen is just beyond the limits of the

<sup>\*</sup> Farey mentions the occurrence of beds of gypsum at Green Hammerton.  $Phil.\ Mag.\ vol.\ xxxix.,\ p.\ 106.$ 

KEUPER. 17

map north of the railway station at Ripon; it is a perfectly perpendicular shaft, having a depth of between 60 and 70 feet. The formation of these pits is, probably, due to the washing away of the friable marks and gypsum below the New Red Sandstone.\*

## KEUPER SANDSTONE AND MARL.

This division of the Trias is not exposed anywhere in the district, and the only evidence of its occurrence is that which can be obtained from boreholes.† There are, however, very few that

throw any light on the subject.

Many years ago a borehole was sunk in search of coal at some distance to the north of Alne Station, which penetrated to a depth of 140 yards, the last part being in grey shale and freestone. At Huby Burn Farm, 66 feet of red marl and gypsum were passed through, At the Union Steam Flour Mills at Easingwold, a considerable thickness of marl was met with. At Raskelf Station, 56 feet of marls and gypsum were pierced before reaching the sandstone below. At Spring House, about half a mile to the south, a great thickness of "freestone rock" was met with. Sandstones also were proved in the borings at Boroughbridge, Brafferton, Newton-on-Ouse, Cattal and Kirk Hammerton. From this evidence it appears that the greater part of the Vale of York is composed of Bunter and Keuper Sandstones, and that the Keuper Marl comes on to the north-east near Huby Burn and Raskelf Station.

In the map to the east white and red sandstones with pebbles of quartz and thin marly partings were found at the North Riding Asylum near York, which probably represent the lower part of the Keuper Sandstone. Red and grey marks and thin white sandstones with ripple marks and pseudomorphous crystals of salt are also seen at several places near the banks of the Derwent, and have been found in borings in the map to the east. These apparently are on the same line of strike as the beds which were penetrated at Easingwold, Raskelf and Huby.

## RHÆTIC BEDS.

Owing to the large amount of Boulder-clay and gravel that covers the flanks of the hills on the eastern side of the Vale of York few opportunities are afforded of studying the junction between the Keuper Marl and the Lias. The existence of the Rhatic Beds is, however, fully proved by sections that are exposed both to the north and also to the east; so that their outcrop to the west and south-east of Easingwold may be inferred

+ Details of these sections are given in the appendix.

<sup>\*</sup> For further account of these pits see papers by the Rev. J. S. Tute, Geol. Mag. vol v., p. 178; Proc. Yorksh. Geol. Soc., vol. v., p. 2; A. C. G. Cameron, Geol. Mag. dec. ii., vol. vi., p. 575, and vol. vii., p. 95; Rep. Brit. Assoc. for 1881, p. 617; Proc. Yorksh. Geol. Soc. vol. vii., p. 342; T. Fairley, Chem. News, vol. xliv., p. 242; Proc Yorksh. Geol. Soc. vol. vii., p. 409; Rep. Brit. Assoc. for 1881, pp. 544 and 601.

TRIAS.

from the general position of the sub-divisions of the Lias which are better seen. The uterop of the Rheetic Bods probably enters the man a short distance to the east of Raskelf Station. an i runs nearly parallel to the railway until it meets the fault running through the town of Easingwold, by which it is thrown my and is met with again on the eastern side of the town, where in a teld alled Paradise Messrs. Tate and Blake mention having found fragments of Rhotic and Keuper rock in the spoilloop from a well From this point the outcrop passes beneath the supermial beds of the flat country, and no indication of the Rhatil Bels is aformed until we reach the slightly rising ground to the south of Stillington. No evidence has been obtained sither as to the thickness or nature of the Rhatic Beds in this immodiate neighbourhood; but in order to give some idea of their character in this district, the following summary, commissi from sections in the country to the north is

nteresting:-	Section of the contract of the tr		2.
		Ft. I	n.
Lins	Brown papery shales with thin wedges of limestone, containing Ostrea liassica, Pleuromya crowcombeia, Protocardia rhatica.  Thin limestone, full of Pleuromya	25	()
	Paper shales, becoming more argillaceous towards the base. Flattened shells or impressions of the same <i>Pleuromya</i> and <i>Cardium</i> , apparently rather dwarfed.	15	()
	Small band of whitish argillaceous lime- stone, about 3 in. thick; no fossils seen White shales, no fossils seen.	7()	()
Rhætic Beds -	Black, crisp, pyritous shales, with flattened Pteria contorta, about 8 ft.  Close-grained, white siliceous bed, Protocardia rhætica, "Isodonta Ewaldi," fragments of Crinoid stems; about 3 in. Black, crisp shales, Pteria contorta very abundant; about 7 ft.  Hard, dark, laminated shale, with fish scales and Pteria contorta resting on close-grained siliceous bed, very full of "Isodonta Ewaldi." The whole is about 9 in. thick, and probably represents the "Bone bed."  Dark shales, 1 ft. 2 in.	17	()
Keuper Marl-	Tea-green marl about	15	0

## CHAPTER V.

#### LIAS.

Lower Lias.—The Lower Lias is but little better exposed than the Keuper and Rhætic beds below, although there are some slight indications of these shales in the neighbourhood of Easingwold; while fragments of the *Pleuromya* limestones are scattered over the fields near Raskelf.

It is, however, only from wells and other artificial excavations

that evidence of these beds can be obtained.

Messrs. Tate and Blake\* record the following fossils, which they assign to the lower part of the Jamesoni Zone, from a drain in the lower part of Easingwold:—Cardiam oxynoti; Pleuromya sp., Belemnites araris, Rhynchonella variabilis, Cacullæa münsteri, Area sp., Ostrea sp., Chemnitzia blainvillei.

Muddle Lias.—In the neighbourhood of Easingwold the Middle Lias is seen in a few obscure sections, chiefly in field-drains and roadside ditches. The sandstones in the lower portion (Margaritatus-beds) are cut through in a small ravine south of Hanover House, and beds in this part of the series were met with in the boreholes for the Waterworks to the north of Howe Hill. The flaggy limestones with Protocardia truncata are also fairly

common and much used for footpaths.

The upper part of the formation (Spinatus Zone) is best seen around Howe Hill, near the Workhouse and near Halfway House; it has been dug out of the foundations of houses in the town of Easingwold, but these were not exposed at the time of our survey. As far as could be judged from the very poor sections visible, its lithological character is similar to what it is on the north-east coast in the neighbourhood of Whitby and Staithes, with the exception that the irony portion is not nearly so strongly (if at all) developed; this southern extremity of the Yorkshire Marlstone appearing to consist only of sandy shale with thin sandstones, which latter are frequently calcareous and very fossiliferous.

From a small stream, shown west of Halfway House, near the

corner of the map, the following species were obtained:-

Protocardia truncata (J. de C. Modiola sp.

Sow).

Rhynchonella tetrahedra (J. (Oxytoma) inæquivalvis (J. Sow).

Rhynchonella rimosa (v. Buch).

In hard shale, a little further on, *Inoceramus dubius* was found.

<sup>\*</sup> Tate and Blake. "The Yorkshire Lias," p. 85.

20 LIAS.

The ironstone of Cleveland does not appear to be of any importance in this neighbourhood, although the chalybeate springs which rise in the alluvium at the foot of the town would intimate that these strata still contain a considerable amount of that mineral; it is probable that its place is taken by less ferruginous beds, for the thickness of the strata between the Upper Lias of Crayke and the base of the sections at Easingwold cannot be much less than 130 feet, which is fully equal to that of the corresponding beds on the coast.

Messrs. Tate and Blake state that 2 feet or more of ironstone was met with in a well at the Workhouse, and that it was likewise reached in a boring for coal near Haverthwaites House, where the rock has a blue oolitic appearance, and contains

Rhynchonella tetrahedra.

Upper Lias.—Two small outliers of the Upper Lias shale occur on the north side of the town of Easingwold, but no sections were observed in them. On the south side of the fault a narrow strip of Upper Lias is found at a higher level at Mount Pleasant, where the lower part of the shales with Inoceramus dubius was seen at the side of the road leading to Easingwold. On the north side they are again thrown up by a cross fault, and crop out around the conspicuous hill upon which Crayke Castle stands.

The thickness of the Upper Lias at this point is about 80 feet; it forms the steep banks below the Castle, and is exposed at the

side of the road.

There are indications that the zones of Dactyliocerus annulatum, Harpocerus serpentinum and Hildocerus bifrons are distinguishable in this neighbourhood, but there are no sections that call for particular notice.\*

<sup>\*</sup> Compare Tate and Blake, "The Yorkshire Lias," pp. 171, 176, 184.

#### CHAPTER VI.

## SUPERFICIAL DEPOSITS.

#### GLACIAL BEDS.

The district described in this memoir comprises a considerable portion of the large central valley of Yorkshire, and in order to give a clear account of the Drift it will be necessary to include a somewhat larger part of this valley than that within the area of

the map.

Throughout the central part of the Vale of York the Triassic rocks are buried beneath a thick covering of Boulder-clay and gravel, varying from about 50 feet to 100 feet in thickness, which gradually ascends the lower flanks of the rising ground on either side. This Drift is sometimes more or less stratified; but a large proportion is piled up in hummocky ridges, the longer axes of which are ranged in a N. and N.W. direction. This is particularly noticeable along that part of the Vale of York which comes within the present district. On this side of the valley the Drift is of two distinct characters, the one containing only fragments derived from the more immediate neighbourhood, while the other includes a great number of rocks that have come from a distance.

The former of these, which appears to be quite a local deposit, occurs in the south-west corner of the map. It contains chiefly fragments and pebbles of rocks that occur in situ at no great

distance.

Pebbles of Carboniferous Limestone, which form so large a percentage of the Drift next to be described, are almost if not entirely absent. This deposit, which is not of great thickness, is bounded on its eastern side by a sharply defined line formed in a general way by the valleys of the Crimple and the Nidd, as far as Ripley, and from thence northwards by the narrow gorge in

which Cayton Gill flows.

East of this line a marked change occurs in the nature of the Drift. It is a much more important deposit than that just mentioned, and has a greater thickness and a more extended range. It contains large numbers of erratic blocks derived from a distance, and is at once distinguished from that to the west by the abundance of Carboniferous Limestone pebbles it contains. This Drift, which is the ordinary Boulder-clay of the country, covers nearly the whole of the area east of the line through Ripley, mentioned above, the solid rock appearing only in a few places where it has been exposed by denudation. It is a portion of that vast accumulation which extends from the country of Durham through the great central valley of Yorkshire. It is confined entirely to this valley and the low hills on its flanks, the

more lofty ranges of the Cleveland and the Hambleton Hills and the Chalk Wolds being entirely free from it. Throughout this region it maintains its usual character of a stiff clay having a dark-blue colour when unweathered; and containing rounded and subangular pebbles and boulders of rock, principally derived from the sandstones and limestone of Carboniferous age as well as of the Oolite and Lias, with a large proportion of granite and igneous rocks, and occasionally a few flints.

Large boulders washed out of these Glacial beds or deposited on their surface have been observed at several places throughout the map; the largest we have seen is a block of Carboniferous Limestone in the bed of the river Skell, near Bishopton Mills, Ripon; this block measures 9 feet in length by  $5\frac{1}{2}$  feet in breadth, with a height of  $3\frac{1}{2}$  feet above the surface of the water; but how far it extended below that level was not determined.

The erratic blocks scattered over this valley are evidently very numerous, as will be seen from the following list showing their distribution, which has been drawn up from the notes of different

observers in this part of the district.\*

Carboniferous Limestone and Sandstone are fairly universal, but special boulders have been noted at Arkendale, Baldersby, Bilborough, Brompton, Carlton Bank, Claro, Coxwold, Harrogate, High Catton, Holtby, Hutton Moor, Kilburn, Kirklington, North Stainley, Sinderby, Stillington, Studley Park, Topcliffe, York, Osmotherley.

Shap Granite at Alne, Baldersby, Bilborough, Claro Hill, Cundall, Holtby, Kilburn, Lindrich, Marton-cum-Grafton, Stillington, Tholthorpe, Thornton-le-Beans, York, Osmotherley.

Lake District Volcanic Rocks at Ainthorpe, Coxwold, High

Catton, Holtby, Thirsk, Osmotherley.

Other volcanic rocks at Aldfield, Bilborough, Brompton, Carlton Bank, Flaxby, Holtby, Hutton Moor, Kilburn, Kirklington, North Otterington, Thirsk, Whorlton, Wighill, Osmotherley.

The western edge of this Drift is sharply marked by a line running from the west side of How Hill along Cayton Gill, by the Nidd at Ripley, to the Crimple near Plumpton, which evidently indicates the extreme limit in this direction of the great glacier that came down the Vale of York. The Drift in this part of its course, that is between Thirsk and York, contains a much larger amount of sand and gravel than it does further north; in fact, the greater proportion of the superficial deposits occurring in this map to the east of the Nidd are of this nature. These gravels are banked up over the rising ground of the Trias and Magnesian Limestone, and form a very striking feature in this part of the country. They occur in a series of more or less continuous ridges running in a S.E. by S. direction roughly parallel with one another and with the general trend of the ground on the west side of the vale. A curious feature connected with these ridges along the western

<sup>\*</sup> Most of these are from the lists in the British Association Reports for 1887, 1888, 1889, 1892, 1893, 1895, 1896, 1897, 1809, 1901, 1902, 1903, 1904.

side is that, while they slope away gradually to the west, they

almost invariably present a steep bank to the east.

These gravels as shown by Professor Kendall, have evidently been piled up along the margin of the great glacier which came down the Vale of York, and at different periods deposited its terminal moraines at York and Escrick.\* By the accumulation of these deposits the water coming off the western uplands was pended back, and forced to seek new channels further to the south. These abandoned channels and the numerous instances of the diversion of the river courses are a very remarkable feature of the district, and will be referred to again later on in treating of the physical structure of the country.

## POST-GLACIAL BEDS.

Much of the sand and gravel forming the higher elevations in the great plain of the Vale of York should probably be classed with the Glacial beds, although they have undoubtedly been much rearranged in Post-glacial times, and it is difficult to separate them from the sands of later date. The same may be said of the laminated clay; this in general character much resembles that of the Tees Valley to the north, which has been considered to be of Glacial age. In this district, however, it joins on without any apparent division to the warp beds of the Humber, which are certainly of quite recent date.

Warp and Lacustrine Clay, Sand and Gravel.—These deposits cover nearly the whole of the surface in the eastern part of the map, forming a portion of the great plain which stretches by York and Selby to the Humber. The whole of this ground, with the exception of Ten Mile Hill, near Tholthorpe, is below the 100 ft. contour line, and if the superficial beds were removed would in most cases be below sea-level. It forms a plain of flat clay land interspersed here and there with low, undulating hills of sand and gravel. These sands and gravels are generally well stratified and without striations.

The warp deposits consist of three divisions:—An upper clay splitting up into little cubical pieces, and often containing small phosphatic nodules (race); a few feet of sand and loam in lenticular masses; and a lower unctuous clay disposed in fine laminæ. At Park House, near Newton, there is a bed of sand below the laminated clay; but it is usual to find this clay in most of the brickyards, where the only good sections are obtained, at the base. These beds have a thickness in some places of 70 feet or more, but it is generally not so much as this.

The lower of these clays only is used for the better kinds of work, such as chimney pots, tiles and best bricks; the upper clay

makes only an inferior kind of brick.

+ See page, 31.

<sup>\*</sup> Proc. Yorksh. Geol. Soc., vol. xii., p. 306; vol. xiii., pp. 15 and 89.

Deposits of peat and clay occur at several places throughout the centre of the map; they occupy depressions in the Boulderclay, and form flat clay land, not unlike the warp country, but on a much smaller scale. In some cases these appear to be due to subsidence over the Permo-Triassic boundary from the washing away of the marl and gypsum near the junction of these formations, as noticed in the chapter on the Trias; but in others, as at Farnham and Goldsborough, they are more probably the result of the blocking-up of old stream courses in Glacial times.

The principal areas of this clay are those which occur west of Farnham, north of Stavelcy, on Arkendale Moor, on Goldsborough Moor, south of Great Ouseburn, west of Cattal Station, and near Hunsingore; but there are a large number of smaller patches throughout the district, a great many of which have been obliterated since the country has been brought under cultivation. Previous to the country being enclosed many of these were mere swamps, but now they are most of them drained by means of deep dykes, and rendered comparatively dry and fit for agricultural purposes.

The following section in beds of this age was opened in 1868, previous to the commencement of our survey of this district; for the particulars of this we are indebted to the Rev. J. S. Tute, of

Markington:—

## Fresh-water deposit near Hollin Hall, Ripon.

1.	Soil -	-	-	-	-	-	-		) ,	10	inches
2.	Yellow	clay	-	~=	-	**	-		} ]	.8	inches.
	Peat	-	-	-	441	-	~	-		3	. 22
	Freshw	ater	shells	(all	of rece	ent s	pecies)	-		1	99
	Marl	-	~	-	-	-		-		6	22
6	Rluigh	elaw.									

Mammalian Gravels.—In the Vale of York, more especially about Overton, are certain sands and gravels in which, from time to time, bones of mammalia have been found. They do not occupy any great extent of surface, but may perhaps extend below the alluvial covering of this great valley. The gravels also at Tollerton, Aldwark, Tholthorpe, and Brafferton are probably of this date, but without further evidence it would be impossible to determine their real age.

The gravels at Overton were formerly extensively used for repairing the roads, and were excavated to a considerable depth. About the year 1818 Mr. James Cook, of York, discovered in these gravels, at a depth of 30 ft., the following mammalian

remains :-

Elephas primigenius (molar tooth). (tusk). Rhinoceros tichorhinus (molar tooth). (portion of tibia). Stag (portion of palmated horn). " (portion of rib). Reindeer (part of an antler). Horse (coronary bone).

Also bones of tiger, elk, and bird.

The bones were found in dark-red clay below 4 feet of rolled gravel, the upper part of the section being chiefly pale-red sand.

Teeth of Mammoth have also been found at Grafton in 1876, near Ripley,\* at Norwood in the Washburn valley, and other places. Deers' horns are recorded from Cowthorpe,† Timble Gill near Fewston and other places in the neighbourhood.

Alluvium and River Terraces.—Along the banks of the Nidd, within four miles of Knaresborough, on either side, and also on the banks of the Skell, near Ripon, there occur several terraces of gravel on both sides of the stream, which were formed when the river flowed at a higher level. The most remarkable of these terraces is that at Ripley, which almost surrounds the boss of Magnesian Limestone standing up at Nidd Rock.

It is a curious fact connected with these terraces that they all occur just above the 100 ft. contour, which is the maximum elevation of the warp clay; and it would appear from this that they were deposited about the same time, and when the lower portions of the Ure and Nidd were under water as far as Ripon and Walshford Bridge respectively, and the warp clay was being deposited in these regions.

posited in these regions.

On either side of the streams, both great and small, there is a modern alluvium—the flat of the present river. This alluvium, which forms the largest speads near the lower part of the Swale and near Ripon, is about 25 feet below the warp flat.

<sup>\*</sup> Proc. Yorksh. Geol. Soc., vol. xii., p. 246. † Phil. Trans., vol. xliv., p. 124.

## CHAPTER VII.

## PHYSICAL STRUCTURE.

On glancing at the map, the first thing that strikes the eye is that all the faults except one are towards the south-west quarter. This arises from two causes: firstly, it is only on this side that the older Carboniferous rocks with their numerous dislocations, and especially the remarkable disturbance at Harrogate, come to the surface; and, secondly, because all the country comprised in the northern and eastern part of the map is obscured by the sands,

clays, and gravels of Glacial and Post-glacial times.

At Harrogate the most interesting portion of the physical geology may be said to be centred; for it is here that we get the eastern end of that great anticline which stretches across the country from Clitheroe in Lancashire to the north of Pendle Hill and Skipton. The direction of this anticline, in the country to the west, is more or less easterly; but within a few miles of Harrogate it suddenly bends towards the north, and in this district has its axis running nearly in a north-east and south-west direction. As it is along this line that the greatest amount of denudation has taken place, and the lowest beds are exposed to view, we cannot do better than commence our detailed account of the position of the rocks by first of all tracing this axis of disturbance

But before going any further it will be necessary to mention the fault which forms the northern boundary between the Millstone Grit and the Harrogate Roadstone Series. This fault, which we will call the great anticlinal fault, has a downthrow to the north; it runs in a north-easterly direction, a short distance north of Shaw Green, along the northern side of Harlow Hill and

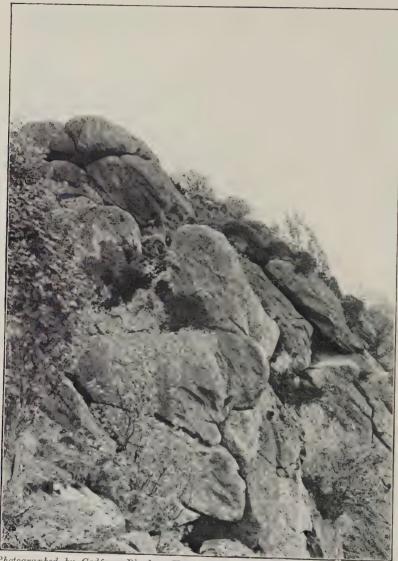
Harrogate, to the Magnesian Limestone near Bilton.

In a small stream, to the south-west of Shaw Green, the Harlow Hill Sandstone is seen first of all dipping to the west at an angle of about 35 degrees, and a little lower down rolling over and dipping at a very high angle to the south-east; on either side of this the encrinital roadstone of Harrogate crops out, dipping to the north and south at angles varying from 20 to 45 degrees. From these dips it is seen that Shaw Green stands upon an anticlinal axis. Now if we follow out the course of these two rocks with the shales both above and below them we shall have described the whole of this series of rocks as far as they are exposed in this part of the country.

To the east of Shaw Green the lower shales crop up and divide the Harlow Hill Sandstone into two portions, the northern of which striking across the low country is lost against the fault near Harlow Carr, while the southern branch (which is much better seen) crosses the country by How Hill and Harlow Hill to

the Bogs Field at Low Harrogate.





Photographed by Godfrey Bingley, Headingley, Leeds.

BIRK CRAGS, NEAR HARROGATE. KINDER SCOUT GRIT DIPPING  $43^{\circ}$  TO THE N.W.

The roadstone, which on the north of the anticline is only seen at one quarry near Shaw Green, to the south has a somewhat extended range, and is well exhibited at several places north of the Crimple Beck as far as Beckwith House, and again near the Cold Bath Road, Low Harrogate, at which latter place it dips to the south-east at about 20 degrees, but soon rolls over, and on the opposite side of the little valley which comes down from the Bogs Field it dips north-west at about 60 degrees into the great anticlinal fault. Between Beckwith House and Cold Bath Road its course is somewhat obscure, and can only be traced by means of the fragments in the fields. A great thickness of shale separates these measures from the lower beds of the Millstone Grit.

The three lowest members of this series, that is the Kinder Scout group, are those which are most affected by the peculiar structure of Harrogate. These coarse and massive grits form, as it were, walls on either side of the anticline, and afford a key to the whole structure of the neighbourhood. To the north of the great anticlinal fault the lowest of these beds is thrown down by that fault, and does not appear at the surface; but to the south the whole series is well exposed in several places, more particularly in the railway cutting and along the hill-side north of Pannal. The uppermost of these beds, which dips towards the Crimple Valley at an angle of from 15 to 20 degrees, forms a well marked and almost unbroken ridge from Almes (Alms) Cliff in the map to the south, through Pannal as far as Starbeck, where it disappears below the Permian rocks.

On the northern side of the anticline the total thickness of this group of rocks does not appear to be so great; but the uppermost member, which dips to the north-west at about 43 degrees, again forms a marked feature to the south of Oak Beek along the

line of the well-known Birk Crags.

These rocks, at their eastern end, bend round towards the fault and at the same time the dip is considerably less, for at Starbeck the top bed of Kinder Scout Grit crops out dipping east at 4 degrees; from this it appears that the anticline of Harrogate dies out to the east, and that there is no great anticlinal ridge of Carboniferous Rocks below the Permian and Trias in this part

of the country.

The Third Grits are not nearly so much disturbed as those we have been describing. To the south of Harrogate the lower or Follifoot beds form the escarpment south of the Crimple, known as Follifoot Moor, where they have a dip to south-east of about 8 degrees. To the north of this they are thrown down by a small east and west fault running through Rudfarlington, but again appear in the valley at Crimple Farm and Rudfarlington Wood, a little north of which they are lost below the unconformable Permian rocks, but again crop up in the bed of the river Nidd above Knaresborough. The Cayton Gill beds, which are not very strikingly exhibited in this part of the country, next succeed and form a band running across the Prospect Tunnel and Rudding Park. To the south-east of this the thick beds of Plumpton Grit

crop out, spreading over the country in the direction of Spofforth and Plumpton with a few inlying patches cropping up from below the Permian at Birkham Wood, Goldsborough Mill and

Knaresborough.

On the northern side of the Harrogate anticline these beds roll over, and on the north side of Oak Beck the Follifoot beds strike north-east across Killinghall Moor and Harrogate End to the Magnesian Limestone at Bilton—the Cayton Gill beds capping the hill at Saltergate. At Killinghall there is a large spread of Plumpton Grit, which, extending eastwards across the Nidd as far as the Limestone near Scriven, is bounded on the west by small faults which again throw up the Lower Third Grit beds on both sides of the river above Ripley Station.

On Scarah Moor there is a second anticline, having its axis parallel to that at Harrogate, which brings up the Follifoot and Cayton Gill beds for a short distance, but they are soon cut off by an east and west fault a little north of Thornton Moor House, and only appear again in the low ground west of Haddockstones Grange and in the valley west of Fountains Abbey.

The Plumpton Grit, north of the river Nidd, occupies a considerable tract of country ranging north-west from Scriven to the banks of the Skell at Fountains. Throughout this district it is much obscured by Boulder-clay and gravel, but its outcrop to

the north and west of Ripley is clearly marked.

The Permian beds, which have a general north-easterly dip of from 2 to 3 degrees, occupy, as we have said before, a band of country running diagonally across the western half of the map. They form a well-marked line capping the grit escarpment from Spofforth to Plumpton, where they are thrown down by an east and west fault, a good section of which was seen in the bank of the river Nidd, and also in the roadside at the same place. The Magnesian Limestone is again seen near Rudfarlington, but a little to the north of this it appears either to thin out or to have been denuded away previous to the deposition of the marl; as in a quarry at Arlington House, red marl rests directly on Millstone Grit rocks, and on the right bank of the Nidd, above Knaresborough High Bridge, a band of limestone appears below the marl, but is too thin to be delineated on the map. About Bilton the limestone again appears to thicken, and it occurs in that neighbourhood in a series of outliers along a north-east and south-west line at Bilton and Scriven. From Knaresborough northwards, the limestone forms a well-marked escarpment by Farnham and South Stainley as far as Markington; from this point northwards its exact base is obscured by a thick deposit of gravel. At Barsneb, South Stainley, Nidd Hall, and Nidd Rock are outlying patches of this limestone; at the first and last of these places the rock is exposed in quarries, but at South Stainlev the only evidence was the fragments in the fields, and at Nidd Hall that derived from a well. At Nidd Rock the limestone appears to be thrown in by a fault, for on the south side of the river the massive red grit of Plumpton forms a cliff of

50 feet or more, and in the bed of the river these beds are seen dipping to the north at a high angle. On the north side of the river nothing but limestone is to be seen until we come to the next railway-cutting to the east, where soft yellow limestone reposes on red sandstone and shales; whether the limestone is here faulted down or was deposited against an old cliff of Millstone Grit there is not sufficient evidence to prove, the river having obliterated the junction.

Although the evidence at this place seems to be in favour of a fault, or at least of an unconformable junction, it is just possible that the apparent irregularity may be caused by a rapid change of dip; for while in the bed of the river, on the southern side, the Millstone Grit is seen to be lying nearly flat, on the northern bank the same beds are seen to dip rapidly to the north and east, and the junction may be conformable and unfaulted as in the diagram given below.



Section across the Nidd at Nidd Rock.

- Millstone Grit.
   Magnesian Limestone.

In the neighbourhood of Goldsborough there are two or three small faults which more or less affect the Permian measures. The principal evidence for these is exposed along the banks of the river, but to the east they soon become obscured by a thick covering of Boulder-clay.



Fig. 6. Fault at Goldsborough Mill. Drawn by J. C. Ward.

a. Red grit. (Millstone Grit.)

b. Soft yellow limestone, with small pebbles from the grit in its lower part, and occasional thin lines of similar pebbles.
c. Yellow earthy thin bedded limestone (above footpath).
d. Earthy limestone.
e. Thin-bedded limestone.
f. Thicker earthy limestone.

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g. Red marl, only a few feet thick, there being a quarry in yellow limestone just behind.

The Triassic beds, which really cover half the superficial area of the map, are, on account of the extensive alluvial deposits of

Permian.

the Vale of York, represented as occupying only a space equa in extent to that of the Permian.

This portion, coloured on the map as Trias, is occupied entirely by the Bunter Sandstone, which has been seen at the few places mentioned above. It is a gently undulating district, thickly covered with Boulder-clay and gravel, and forming no good geological boundaries. On the western side its base can be only traced approximately by the aid of wells, and on the eastern side it gradually sinks below the great plain of central Yorkshire leaving no clue, except from a few boreholes, as to the nature of the strata between this and Easingwold, where the fossiliferous beds of the Lias begin to rise on the flanks of the Oolitic outlier at Crayke.

Thus all the faults, with the exception of those near Goldsborough, appear to be more or less connected, and to owe their origin to the great pre-Permian disturbance in the western part of the map; for at Harrogate all the faults of any note are more or less parallel to that axis, and this is again the case on Scarah Moor; but in the intervening area between these to the west of Ripley, where the beds are somewhat arched along an axis striking north and south, the faults also range in that direction.

From the foregoing remarks it is seen that the area included in the map may be divided into three districts, each having an aspect peculiar to itself. In the south-west corner the Millstone Grit rocks, when not obscured by Drift, form the usual bold features and deep valleys peculiar to this formation, although not in so marked a degree as these rocks do further to the west, the effect being somewhat lessened by the high angles at which the rocks lie near Harrogate. The series of rocks beneath the Millstone Grit attain a height of 600 feet above the sea at Harlow Hill, but the Millstone Grit rocks themselves do not rise to more than an average height of 400 feet.

The Permian and Triassic rocks, or, more properly speaking, the Magnesian Limestone and Bunter Sandstone, as it is only these which really affect the physical character of the country, form the central district. These rocks, which together occupy an area equal to more than half the map, usually form a good escarpment to the west, while to the east they gradually slope away below the alluvial covering of the Vale of York. This district, which is somewhat elevated, and in the extreme north, near Fountains Abbey, where it is thickly covered with Drift, attains a height of nearly 600 feet, does not rise to more than 200 feet above the sea in the southern portion of the map.

Almost the entire surface of these rocks is covered by gravels and sands of Glacial and Post-glacial age, which have very much altered its physical aspect; so that instead of a continuous and uniform slope to the east, we get a series of mounds and hills which form a generally undulating and park-like country.

To the east of this is the great alluvial plain stretching north from York. This district, which, when viewed from the higher

ground on either side, appears to be a dead-flat, is in reality somewhat undulating. The Warp Clay, which covers the larger part of this ground, forms an almost perfectly level plain, having an average height of about 50 feet above the present sea-level, while the sand and gravel with which it is associated forms gently undulating ground, attaining at Tholthorpe a height of 75 feet above this plain.

Before closing this portion of our subject it will be as well to say a few words about the general denudation of the valleys and the courses in which the present rivers flow.

The general drainage of the country is to the east, as might be naturally inferred from the dip of the rocks; but it is curious to observe that all the rivers to the west of the Magnesian Limestone escarpment have cut their way through that formation, and form a series of steep valleys or gorges, which contain some of the prettiest scenery in the district; thus the Skell at Studley, the Beck at South Stainley, the Nidd at Knaresborough, the Crimple at Spofforth, and even the little stream near Walkingham Hall have each a separate channel through the Permian escarpment.

These valleys however are not in many cases the original courses of the streams, which have been diverted and forced into new channels by the accumulation of Glacial detritus. Thus the Laver, which flows to the north of Grantley and Aldfield, at one time turned south along the narrow valley between these places now occupied by the Horsley Dike, and joined the Skell below Grantley Hall.

The Skell appears to have twice altered its course. In its upper part it once occupied the streamless defile east of Sawley Hall known as the Dean, and joining the Hebden Beck flowed southwards. Lower down, after traversing the grounds of Fountains Abbey, it flowed north and east across Studley Park; this valley becoming blocked by Glacial beds, it cut the present picturesque channel through the Magnesian Limestone in Mackershaw Woods to the east of the great lake.

Another instance is the case of the Markington Beck. This stream coming down from Haddockstones now flows eastwards by Markington and South Stainley into the Ure at Newby; this does not appear to have been its original course, for at Dole Bank and thence southwards to Ripley, there is a well-marked gorge, which is in some places fully as wide and deep as that of the Nidd near Knaresborough, and considerably larger than that in which the present stream flows. This valley is now drained only by the diminutive stream of Cayton Gill, which joins the Nidd below Ripley; but it is evident that a much larger stream must originally have cut the streamless defile at Dole Bank, where it is interesting to observe the abrupt turn the Markington Beck makes to avoid this valley.

The River Nidd furnishes a further instance of the alteration of the stream courses. This river, which now flows through the evidently Post-glacial gorge past Bilton and Knaresborough,

originally passed to the north of Nidd Rock and followed the more open valley by Nidd Hall and Brearton, to the gap in the

Magnesium Limestone at Staveley.

Thus, all these streams have been diverted from their original channels, which are now marked by the almost streamless defiles at Grantley, Sawley, and Dole Bank; and which appears to have been the course of a southward flowing stream that intercepted all the drainage coming from the west.

#### CHAPTER VIII.

## ECONOMIC GEOLOGY.

Reference to the economic value of the rocks and minerals has been incidentally made in the previous pages. A further account is here given with special regard to these points.

Mineral Waters.—Of all the products derived from the rocks in this district the most important are the mineral waters to which the town of Harrogate owes its origin and success. These springs are confined almost entirely to the south-west corner of the map, although mineral springs are found at a few other places in the neighbourhood, and in one or two instances have also been reached by borings. They are situated on the axes of the great folds which have taken place in the older rocks, or in close proximity to these anticlines, and there is but little doubt that they owe their origin to this cause. The history and general occurrence of the Harrogate waters is treated of in the next chapter, so that it is not necessary to go further into the subject here.\* But besides the eighty or so springs which rise from the Harrogate anticline there are a few other mineral springs rising in the neighbourhood which may be briefly noticed.

A sulphur spring issues at Aldfield, in the valley of the Skell above Fountains Abbey on the axis of an anticline parallel to that at Harrogate. This spring, being somewhat difficult of access, has not been much noted. An old analysis of this water was made by Dr. W. Brunton in 1806, which is quoted by Dr. G.

Cayley. †

		Gra	ains pe	r gallon.
Carbonate of lime -	~	**	-	12.5
Carbonate of magnesia	***	40	-	3.5
Sulphate of magnesia	200	-	-	5.0
Muriate of magnesia	-	**	-	96.0
Muriate of soda -	~		-	2080
		(	Cubic I	Inches.
Carbonic acid gas			-	6
Azotic gas	-		-	4
Sulphuretted hydro	gen -	. ,	- 2	1
Temperature $54^{\circ}$	F.			

<sup>\*</sup> See page 42. † Dr. G. Cayley, "Synoptical Tables," 1809. 10514

The following more modern analysis of this water has been supplied by the Manager of the Spa Hydro at Ripon.

Extract from the Report of Professor (now Sir Thomas)

· Stevenson, Guy's Hospital, London.

Specific gravity of the water at  $60^{\circ}$  F. 1.0032.

Temperature of spring 53° F. (air temp. 70° F.).

				In grains	In parts
				per gallon.	per 100,000.
Strontium Sulphate	_	-	-	0.76	1.09
Calcium Sulphate	-	-	-	22.10	31.57
Calcium Carbonate	-	-	-	9.84	14.06
Calcium Chloride	-	-	-	6.48	9.26
Magnesium Chloride	-	-	-	30.54	43.62
Magnesium Bromide	***	-	-	0.13	0.19
Magnesium Iodide	-	~	~	Traces	Traces
Potassium Chloride	-	-	~	4.37	6.24
Sodium Sulphate	-	-		5.51	7.73
Sodium Chloride	-	-		195.27	278.95
Lithium Chloride	-	-	-	0.87	124
Ferrous Phosphate	-	-		0.06	0.09
Ammonium Carbonate	9	-	-	0.23	0.33
Silica	-	~	-	1.10	1.57
				277.16	395.94
Sulphuretted Hydrogra	en	-	-	0.77	1.10

Or in cubic inches per gallon at 60° F. 2.1.

N.B.—The Sulphuretted Hydrogen was all, or nearly all, in the free or uncombined state.

## Examination for Organic Purity.

					In parts
				per gallon.	per 100,000.
Saline Ammonia	-	-	-	0.081	0.1157
Albuminoid Ammonia		-	-	0.001	0.0014

A spring of considerable celebrity is the Dropping well at Knaresborough. This, which is an ordinary petrifying spring, has drawn much attention from the picturesque manner in which it spreads over a limestone crag, covering it with a thick tufaceous deposit. This spring, which has been mentioned or described by many authors writing on the district, issues from the Magnesian Limestone, but whether the water is thrown out from the junction of this limestone with the Carboniferous rocks below is uncertain\*. The following analysis of this water is given by B. A. Burrell:—†

† Journ. Chem. Soc. 1896, vol. lxix., p. 536. Proc. Yorksh. Geol. Soc. 1897, vol. xiii., pt. ii. p. 141.

<sup>\*</sup> A view of this spring, published by J. Smith in 1747, is in the British Museum Library, K. 44 (38,39).



Photographed by Godfrey Bingley, Headingley, Leeds.

The Dropping Well, Knaresborough. Crag of Magnesian

THE DROPPING WELL, KNARESBOROUGH. CRAG OF MAGNESIAN LIMESTONE COVERED WITH TUFA.



# Dissolved Saline Constituents in grains per gallon:—

Calcium Sulphate -	84	_	_	114:373
Calcium Carbonate	~	_	_	25.480
Calcium Silicate -	***	_	-	1.442
Strontium Sulphate		-		0.672
Magnesium Sulphate	_	~		17:003
Ferrous Carbonate -	_	***	_	0.154
Manganese Carbonate	_	_	~	0.084
Sodium Chloride -	_	_	***	2.765
Potassium Chloride	-	_		0.462
Alumina			_	trace
Phosphoric Acid -	_	_		trace
Troophorio Hold			_	orace
				169:435

162.435

Fixed residue by evaporation dried at 170 to 180° C. - 162·324. Specific gravity at 15·5° C. - - - - 10024,

Gases dissolved by the water in cubic inches per gallon, and measured at 15.5° C., and 760 m.m.

Carbon Di-oxide	-		-	~	4.74
Oxygen	-	_	-		1.29
Nitrogen -	-	-	. –	-	5.02
					11:05

Another spring which has been alluded to by many of the older writers is that of St. Mungo. This does not appear, however, to be a spring of very marked mineral character, and is said to owe its reputation chiefly to its great coldness. There is some confusion as to the situation of this spring. There is little doubt, however, that it is the spring near Copgrove, which is marked on the Ordnance map, but many of the older writers place it about half a mile west of the Old Sulphur Well.\*

Building Stone.—The best building stone of the district is undoubtedly the Millstone Grit, which has been quarried at numerous places along the western side of the map. The Magnesian Limestone also makes a good building stone, but its quality in this district does not appear to equal what it is further south, where it has been used in many important buildings, the Houses of Parliament, Westminster Hall, York Minster, Beverley Minster and many other structures being built of stone procured

<sup>\*</sup> St. Mungo's (or Mongah's) well was visited by Thoresby in 1682; Rev. J. Hunter, "Diary of Thoresby," vol. i., pp. 86, 234, 424. A full account of this well is given by D. H. Atkinson, "Ralph Thoresby the Topographer, his town and times," 1885, p. 114, and add. notes. The earliest description of this spring is by Dr. Clayton in 1697. See Sir J. Floyer—"An Enquiry into the Right Use of . . . . Baths in England, pp. 113-117. Dr. Garnett places this well half a mile to the east of the sulphur wells at Harrogate. S. Lewis in his "Topographical Dictionary," 1849, says it was in Bilton Park, while in Hargoves "History of Knaresborough," &c., it is given as one mile west of the Queen's Head.

from quarries in this formation.\* The stone used by the Romans in the building of Isurium (Aldborough) is supposed to have come from the Magnesian Limestone in the neighbourhood of Staveley.+

It is, however, too easily affected by the weather to form a lasting material for outside work. On exposure to frost the mass of the rock flakes off and crumbles away, leaving the sparry veins standing out as ridges. For inside work it is well adapted, as the stone can be easily worked and is of a good colour.

The only other rock that can be used in building construction is the Bunter Sandstone, but this is generally far too soft to be available for the purpose. It has however been used to some extent in the neighbourhood of Ripon and Boroughbridge.

About Easingwold the flaggy beds of the Middle Lias have been used for paving footpaths, but the outcrop is too limited for the practice to become very general.

Lime.—The Magnesian Limestone has been extensively quarried for burning lime. The rock has been worked at many places, particularly near Knaresborough and to the south of Ripon, but the most extensive quarries at the present time are those at Wormald Green. The rock here is a grey magnesian limestone, which when calcined gives the following analysis.

## Analysis of Lime.—Wormald Green.

V					
Calcium oxide	-	-	-		72.71
Magnesia	-	-	-	-	19.26
Oxide of Iron and	d Alur	nina	-		1.55
Silica	-	-	-	-	1.76
Carbon di-oxide	***	-	~	•	3.05
Sulphuric anhyd	ride -	-	-	-	.97
Combined water,	&c	-	-	-	.70
				-	
					00.001

100.00

This varies a little, having sometimes a larger percentage of calcium oxide and less magnesia.

Some of the quarries in this neighbourhood consist of hard grey and blue beds, which are very valuable for agricultural lime, owing to the large percentage of calcium oxide that they contain; whilst the whiter beds containing a large percentage of magnesia make a valuable building lime.

At Well, to the north of Ripon, the rock is more uniform and

<sup>\*</sup> York Minster was partly built of stone from Jackdaw Crag, Thorpe Arch; and Beverley Minster of that from Smawse Quarry, Bramham Moor.

<sup>+</sup> J. J. Sheahan and T. Whellan, "History and Topography of the City of York," vol. iii., p. 182.

<sup>‡</sup> For the particulars and analysis of the lime at Wormald Green and Well we are indebted to Mr. F. Hymas, of Burton Leonard.

blue in colour, forming a good road-metal and excellent for agricultural lime, of which the following is an analysis.

## Analysis of Lime at Well.

Calcium		***	-	wa	_	-	86.51
Magnesia		~	v 100		-		1.39
Oxide of	Iron a	and A	lumii	na			2.77
Silica-	_	_	_	_	-	_	3.73
Carbonic	anhy	dride	, coml	bined	wat	er,	
&c.	-	-	-		-	-	5.60
							100:00

As the chemical composition of the Magnesian Limestone appears to vary somewhat throughout this area, it may be interesting to compare the analysis of the unburnt limestone in the south of the map with that in the north. The analyses, by Mr. Holme, given below are quoted from Professor Sedgwick's paper on the Permian Series.\*

# Earthy variety of Magnesian Limestone from Ripon.

Carbonate of lime		-	-	-	71.125
Carbonate of magnesia -	~	-	~	***	25.625
Red oxide of iron and alumina	-		-	-	1.750
Silica, a trace of, and water -	-	-	-	-	1.500
					100 grs.

# Earthy variety of Magnesian Limestone from Knaresborough.

-		-		72.00
-	-		-	25.50
	-		**	1.50
-	-	-	-	1.50
				100 gr
	-			

The lime from the Magnesian Limestone is said to produce a mortar which sets with extreme hardness; but for agricultural purposes it is too caustic to allow a large amount to be used at once.

Coal.—Two or three thin coal seams occur in the Millstone Grit rocks. These were many years ago worked at Bilton, and used for burning the Magnesian Limestone of the neighbourhood. The upper seam was 3ft. 2ins. thick, while the lower was 2ft. 9ins. but thinned away towards the north. The measures are much disturbed here by the Harrogate anticline, so there would probably be much difficulty in working these thin seams.

These seams have also been worked in several places to the north and west of Harrogate, but they do not appear to have

<sup>\*</sup> Trans. Geol. Soc., ser. 2, vol. iii. p. 87.

been of much importance. Winch mentions coal being found at Easingwold,\* but this probably refers to the coals which were obtained from the Oolites further to the north.+

Sand and Gravel.—Thick deposits of sand and gravel occur very generally throughout the district, especially over the central area, and in the low ground east of the River Ouse. The coarse glacial gravels, which occur at Grafton and elsewhere, were formerly extensively worked for road material; but since the introduction of the igneous rocks for this purpose most of these pits have been closed.

There are also some extensive deposits of river gravel along the banks of the Nidd below Knaresborough which have been

worked from time to time.

At the base of the Magnesian Limestone we have noticed certain sands, which in the neighbourhood of the Coalfield were used for moulding purposes. The outcrop of these sands in the present district appears to be very limited; but if they should prove to be of value it is very probable that further exploration might reveal a more extended development of them.

Gypsum.—Gypsum occurs in the Magnesian Limestone at Bilton, in the Trias at Green Hammerton, and in the marls on the banks of the Ure at Ripon. It does not appear, however, to have been worked to any extent, and the beds are probably too thin to be of much commercial value.

Celestite and Strontianite.—Sulphate of Strontium occurs on the banks of the Nidd at Biltons and also at Green Hammerton. Carbonate of Strontium is stated to have been found at Merryfield near Pateley Bridge.

Copper-ore.—Green Carbonate of Copper is said to have been obtained in considerable quantity at Farnham, north of Knaresborough, "by means of galleries worked through the limestone," and at Newton Kyme, near Tadcaster.\*\*

Water Supply.—The principal water-bearing rocks of the district are the Millstone Grit, the Magnesian Limestone and the

\*Phil. Mag., vol. xlvii., p. 100. + See Mem. Geol. Survey. "The Geology of the Country around Northallerton and Thirsk," p. 37.

ton and Thirsk," p. 37.

‡ The occurrence of alabaster is noted by Dr. French near Harrogate, who also states that "glittering sand with some gold" was found there.—"The Yorkshire Spaw," 1652.

§ J. Sowerby, "British Mineralogy," 1817, vol. v., p. 75, tab. 444. Dr. Murray, Trans. Geol. Soc., 1817, vol. iv., p. 445. Crystals of Celestine from here are in York Museum.—Ann. Rep. Yorkshire Phil. Soc. for 1887, p. 34.

¶ J. Farey, Phil. Mag. 1812, vol. xxxix., p. 106.

∥ Edin. Phil. Journ. 1825, vol. xii., p. 178. This is stated by R. H. Davis to be the double Carbonate of Baryta and Strontia.—Chem. News, vol. xiii. p. 303.

p. 303. \*\* W. Marshall, Trans. Geol. Soc., ser. 2, vol. ii., p. 140.

Triassic Sandstones. The Millstone Grit, however, must be regarded as the most important source of supply; but its outcrop in the immediate district covers such a limited area, and the rock itself has been subjected to such frequent folding and faulting, that the amount of water available from these rocks cannot be expected to be very great. For these reasons the larger towns, Harrogate, Knaresborough and Ripon, have been compelled to seek their supply from a greater distance, and to construct reservoirs in the upland valleys to the west.

At the present time Harrogate derives its supply from reservoirs at Scargill and Beaver Dyke in the drainage area of the Oak Beck, which collect the water falling on the moorlands to the west; but these will be supplemented by further works that are being undertaken at Roundhills in the valley of the Pott Beck, a tributary of the Burn which flows into the River Ure at Masham.

The following particulars as to the Harrogate water supply have been furnished by Mr. F. J. Dixon.

The works were established in 1846 and purchased by the Local Authority in 1898. The Storage Reservoirs are:—

			Capacity		Watershed
			in gallons.		in acres.
Scargill	-	-	200,000,000		1,100
Upper Beaver Dyke	-	-	30,000,000)		1,260
Lower Beaver Dyke	-	-	118,000,000	* * *	1,200
Ten Acres	-	-	35,000,000		410
Roundhills	-	-	575,000,000	* * *	3,000

The average annual Rainfall over the Drainage Area is 28:58 inches (period 11 years). Calculated 30 years mean 30:20 inches. The area supplied by these works includes the towns of Harrogate and Knaresborough together with the parishes of Pannal, Bilton, Starbeck, Scriven and Knaresborough.

The analyses of two of these waters are as follows. The samples contain in grains per gallon (parts per 70,000), after filtration:—

COHOURT HE	T COLLEGE	0 m		, Louis	L.		,	/)		
									Scargill.	Beaver Dyke.
Chlorides equ	al to cor	nmon	salt	-	den .	in .	-	~	1.84	1.84
Calcium sulp	hate	-		-	-	-	-	-	4.55	4.08
Calcium and	Magnesi	u <mark>m</mark> ca	rbona	ites,	&c.,	-	~ 1	-	trace	0.09
Volatile and	organic	matte	r	- 1	-		-	-	0.02	1.80
Total dissolve		-	-	**	-	-	100	-	6.44	7.81
Chlorine .		_	-	-	-	***		-	1.15	1.15
Ammonia .		-	-	-	-	-			0.008	0.002
Organic amm	onia	-		-	-		at .	-	0.003	0.007
Oxygen requi	red to ox	idise	organ	ic m	atter	in fo	ar het	ırs	0.003	0.148
7781			Ŭ						Degre	ees.
Hardness befo	ore boilir	g	_	_	_	-	60	-	2.7	3.1
Hardness afte				-		-		all	2.5	3.0

The Ripon supply is obtained from a reservoir having a capacity of about 88,000,000 gallons on Lumley Moor,\*about five miles

<sup>\*</sup> Lungley on Ordnance Map.

west of the town. The following analysis by Mr. A. H. Allen (September 1899) has been communicated by Mr. J. W. Kirkley:-

Grains Grains per gallon. per gallon. 1.04 5.74, which lost on ignition -Total solid matter 0'90, equal to Sodium chloride 1.48 Chlorine Nitrogen in oxidised forms None, equal to Nitric acid (anhydrous) None Degrees. 6.0 Hardness -Parts

per million Oxygen absorbed 0.95 Free ammonia -0.04 Albuminoid ammonia

The water supply of Easingwold is derived from the Middle Lias; while excellent springs of water also issue from the base of the Oolites in that neighbourhood. The following analyses show the character of some of the waters derived from these sources :-

Communicated by Messrs. Fairbank and Sons, Driffield. Analyses of three samples of water by J. Baynes. March 1890.

	A A spring opposite Banks Farm.	B Hanover Farm.	C Oulston Spring	Common relations
Total solid residue -	22.40	22.00	26*25	Grains per gallon.
Chlorine	1.10	1.23	•90	De-t ili
Free Ammonia -	.0214	None	None	Parts per million.
Albuminoid Ammonia	.0230	.008	.014	

A [Probably from the Middle Lias, but through a thin covering of Boulder

C. [Issues from the base of the Lower Oolites.] A water of extraordinary

degree of purity.

Both the Magnesian Limestone and the Triassic Sandstones, from their open porous character, collect a large amount of water, which has been utilised in some cases; but, owing to the quantity of gypsum associated with these rocks, it is liable to be excessively hard, and quite unfit for domestic purposes.

Besides these rocks a considerable amount of water may be obtained from the Glacial gravels, which are scattered over a large part of the district, especially along the range of rising ground in the centre of the map, and over the undulating hills

clay,] A water of a high degree of purity,
B. From a boring in Bog Field about 300 yards east of Banks Farm.
[From the Boulder clay overlying Lias, probably from the Middle Lias.] A water of extraordinary degree of purity. Issues from three 3 in. boreholes at the rate of 54,648 gallons per 24 hours (April 1891). This is the water selected for the supply of the town of Easingwold.

in the Vale of York. Its occurrence, however, is very uncertain, and as in the neighbourhood of villages it is liable to contamina-

tion, it is not a source to be generally recommended.

The only remaining source of water supply is that to be derived from the rivers. This in the case of York is drawn from the River Ouse, a short distance below the drainage area of this district. Two analyses of this water are added for comparison:—

# Analysis of Filtered Water. From a Report of the York Waterworks Company 1903

Trom a report of th	IC TOIR	44 00	DOL MOL	NO C	սուր	aшy,	1900.
					_	,	Parts
						per	100,000.
Silica	**		•	-	-	-	•32
Oxide of iron and alu	ımina	_	-	-	_	_	.72
Carbonate of lime -	- 1	_	_	-	_	42	11.27
Sulphate of lime -	1.2	-	-	_	_	_	2.57
Nitrate of lime -	-	_	-	_	_	_	*21
Sulphate of magnesia		_	_	_		_	1.82
Chloride of sodium	_	**	-		_	_	2.58
Organic matter and	loss			_	-	_	1.61
			Total	solid	l res	idue	21.10
							Parts
						per	100,000.
Chlorine		-	-	-	-		1.57
Free ammonia -	av .	_		-	-	-	0.000
A 11 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1							0.004

								POI	100,000.
Chlorine	-	-		-	-	-	44	-	1.57
Free ammon	ia	-	er .	-		-	-	-	0.000
Albuminoid	amm	onia	-	~	-	-		un	0.004
Oxygen abso	rbed	in 1	5 min	utes a	t 80°	F.	-	-	0.36
Oxygen abso	rbed	in 4	hours	at 80	°F.	-	-		0.40
Nitrogen in							-	-	0.024
9									Degrees
									Clarke.
Temporary h	ardn	ess	-	_		_		_	7.2
Permanent	**				20	_	-	-	1.5
	"								
				To	tal h	ardn	ess	-	8.7

Analysis of the water supplied by the York Waterworks Company, by Mr. Patterson, December, 1876.

## Local Government Board Report, 1885. No. 4.

Local Governmen	U DO	aru r	cpor	0, 100	0.	TIO.	T.
						(	Frains
						per	gallon.
Total solid matter drie	d at	212°	F.	-	-	-	15.700
Chloride of sodium	-	91	**	-			1.500
Ammonia	~	**	-	-	***	-	trace
Albuminoid ammonia	-	-	-	-	-	-	0.002
Nitrogen existing as n	itrite	s and	nitr	ates	-	-	0.059
Oxygen required to ox	idise	orga	nic m	atter	-		0.125
1							Degrees.
Hardness before boiling	ıg	•~			-	-	12.1
Hardness after boiling		hour	-		-	-	7.4

The water was very clear, containing a very minute trace of matter in suspension. It had a slight brownish-yellow tinge of colour, arising from the presence of peaty matter; the amount was so slight as to be unobjectionable. The water is quite free from odour, and it is very pleasant to taste. It is very free from nitrogenous organic impurities, and is a wholesome and pleasant drinking water.

(Signed) JOHN PATTERSON.

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#### CHAPTER IX.

## HISTORY AND ORIGIN OF THE HARROGATE SPRINGS.

Of the early history of Harrogate very little is known. The place is first mentioned in the Stuteville charter about 1200, where it is designated Harelow.\* With regard to the etymology of the name there appears to be some difference of opinion amongst the older writers. Hargrove states that it was originally called Herrigate, from the frequent invasion of the Danes and other nations. + Hunter and Lewis give the old name as Heywragate, from its situation on the road from Knaresborough to Heywray (Haverah Park). The great forest of Knaresborough which extended from that town to the valley of the Wharfe about Bolton Abbey included the area where Harrogate now stands. In the reign of Elizabeth much of the timber was cut down for smelting the iron ores of the neighbourhood.\(\frac{1}{2}\) Harrogate, however, is rarely mentioned before the discovery of the first spring in the 16th century. This was in 1571, when Capt. Slingsby, on his return from Westphalia, discovered the Tewit Well, or English Spaw, noticing that the water was very similar to what he had seen in that country. Soon after its discovery Dr. Bright wrote on its virtues and uses, and in 1626 Dr. Dean published his Spadacrene Anglica, which is the earliest general account of these waters. The next spring to be discovered was the Old Sulphur Well, and this was followed by the Old Spaw, or Sweet Spaw, afterwards called John's Well.\*\* The dates, however, of these discoveries are uncertain, but as the Old Sulphur Well is mentioned by Dean, it must have been known previous to 1626, and was not discovered subsequently to that date, as has been stated by some authors. The Old Spa is claimed to have

and in those by J. Thorpe and T. Inchbold, it is stated that the Old Spaw was

and in those by 3. Intorpe and 1. Inchrond, it is stated that the Old Spaw was the first well discovered by Slingsby, but this is an error.

|| Dr. W. Buchan states that the Harrogate Waters were first brought into repute by Dr. Daltry, of York. "Cautions concerning Cold Bathing, and drinking Mineral Waters," p. 16. 1786.

\*\* The name John's Well was derived from John Hardestie, the attendant.

It is often erroneously called St. John's Well.

<sup>\*</sup>By this charter a portion of the Forest of Knaresborough, including Harrogate, was granted by William de Stuteville to the Plumptons.

† Here in Saxon signifying an army. "History of the Castle, Town and Forest of Knaresborough." By E. Hargrove, 1775, p. 41.

‡ "A Treatise on Mineral Waters of Harrogate and its Vicinity," By Dr. A. Hunter, 1830. "Topographical Dictionary of England." By S. Lewis, 1849.

§ These furnaces were at the neighbouring village of Kirby Overblow, which in an old MS. of the time of Edward I. was known by the name of Kirby-Ore-Blowers. Ralph Thoresby, "Ducat Leodiensis, 1715, p. 166.

¶ In Hargrove's History of Knaresborough and the Guides published in 1840, and in those by J. Thorpe and T. Inchbold, it is stated that the Old Spaw was

been discovered by Dr. Stanhope in 1631. During the 17th century the importance of the springs became more known, and

the accommodation for visitors had to be increased.\*

During the greater part of the 18th century the country around Harrogate was very wild and uncultivated, but in 1770 an Act was passed for the enclosure of the Forest of Knaresborough, and the value of the mineral waters was so prized that 200 acres of land, including many of the springs, were reserved to be free and unenclosed, forming the common now known as the Stray.

The Alum Well, in the Bogs Field, was discovered in 1733,

being subsequently lost, but re-discovered again in 1870.

The spring in the Crescent Gardens was discovered by chance

in digging for water in 1783.

In the early part of the 19th century the mineral waters of Harrogate became more appreciated, and various other springs were discovered from time to time—the Cheltenham (Chloride of Iron) spring in 1819<sup>+</sup>, those in the Montpellier Gardens in 1822. and those at Harlow Car about 1840, while since that date many others have been added to the list; so that at the present time there are said to be 80 known springs, differing in strength and quality, the Bogs Field alone containing 36 separate springs. The existence of so many springs, no two of which are alike. issuing side by side, is unique, and has given rise to various opinions from geologists and others as to the probable causes of this remarkable occurrence. Dr. Garnett, William Smith, Edwin Lee, Professors Phillips, Turner, Williamson, Daniel and Johnstone, William West, Nicholas Brown, Dr. Murcay, Dr. Clanny, John Buddle, Dr. Oliver, and others have all expressed opinions as to the origin and cause of these springs. † Amongst these there is a general agreement, although they differ on minor points, that the majority of the springs are deep-seated, that the water is obtained from a distance, and that the cause of the great variety of the springs and of their issuing at this particular spot is due to the peculiar geological structure of the neighbourhood.

A full account of the geology of the district has already been given in the previous pages of this memoir, so that it will not be necessary to recapitulate more of the account of the general structure of the rocks than so far as it affects the origin of these springs. Harrogate, as we have noticed, is situated at the eastern termination of the great anticline which traverses this part of Yorkshire; and there can be but little doubt that it is due to this fact that mineral springs occur here in such abundance, and of so varied a character. The strata which are bent up by this anticline are the lower members of the Carboniferous series—the Millstone Grit, the shales and sandstones associated with the Harrogate

<sup>\*</sup> The first inn, called the Queen's Head, now the Queen's Hotel, was built in 1687.

+ This spring was discovered by boring. R. H. Davis, Chem. Soc. Journ.,

<sup>1873,</sup> pp. 1089-92.

† These statements and opinions are published by W. Grainge in his Geology of Harrogate, 1864.

Roadstone, and the Mountain or Scar Limestone; these three principal divisions are further sub-divided into their separate beds of sandstone, shale, limestone or chert, some of which have very distinctive characters, which they retain throughout large areas. The Millstone Grit of this district is for the most part a series of some half-a-dozen beds of coarse grits, separated from each other by thick beds of shale. A very good section of these beds is exposed along the railway cutting from Harrogate to the south; but the best idea of the nature of the grits themselves is obtained from Plumpton Rocks, Birk Crag, and Almes Cliff Rocks. It is these rocks which, stretching over the moorlands to the west, form some of the grandest and wildest of the Yorkshire scenery.

The series of rocks, which include the beds that we have called the Harrogate Roadstone, consist in the upper part of dark blue shales, these being succeeded by thin bands of earthy limestone and chert containing Encrinites. This latter is a hard siliceous rock, with a conchoidal and hackly fracture, but it is very frequently much decomposed, having the calcareous portion dissolved out, when it resembles pumice-stone in texture. It is extensively quarried in the neighbourhood for mending roads. These cherty beds rest on dark blue shales containing sulphur, iron, and probably several other chemical substances, and below these again there is a thin rubbly sandstone also containing sulphur and iron, resting on other beds of shale, which are the lowest strata visible in the district.

The precise equivalents of these measures to the west are rendered somewhat doubtful by the thinning out of one or more of the beds to the west of Harrogate; for in the country west of the Millstone Grit area, where we should expect to find the same measures cropping up, there are no rocks that exactly correspond to those at Harrogate. The reason of this is, either that these measures gradually become thinner until they finally disappear in their passage to the west, or that they are so altered in their general composition as to be no longer recognisable as the same beds. When such changes as these take place in an isolated area like Harrogate, it is extremely difficult to identify the beds or to say exactly to what part of a formation they belong. In this case the only clue to their correlation with the strata of a known area is the lithological character of the rocks, and even this is not always to be depended upon. But we shall not be far wrong in considering these beds at Harrogate a portion of the Pendleside Series, for it is evident they lie below the Millstone Grit, while at the same time they must be above the Scar Limestone.

The strata which occupy the district to the west of Harrogate and south of Settle are thrown into a series of nearly parallel anticlinal and synclinal folds, ranging E.N.E. or N.E., of which the most remarkable are those in the neighbourhood of Bolland and Clitheroe, by which the Scar Limestone and the beds of the Pendleside Series are frequently brought to the surface, and extend in long lines across the country to Skipton and Appletreewick,

whence the direction is nearly east towards Harrogate and Ripon. Along these lines the higher and middle members of the Millstone Grit rocks are wanting, and the lowest beds of Millstone Grit crop out, enclosing elongated patches of Lower Carboniferous rocks. If it were not for these anticlines, the Millstone Grit might be separated from the rocks below by a nearly north and south line ranging along the western side of Great Whernside, Rylstone Fell, and Skipton; then in a general way, with the exception of some of the higher hills which are capped with Millstone Grit, all the country lying west of this line would be composed of Lower Carboniferous rocks, while that to the east would have the various members of the Millstone Grit; as it is, this regularity is broken by the disturbances mentioned above, so that the base line of the Grit retreats considerably further to the east in the neighbourhood of Greenhow Hill and Skipton.

On the northern of these two lines of disturbance the sulphur well at Aldfield is situated, while the southern anticline upon which Harrogate stands is a continuation of that at Skipton, although its direction in the neighbourhood of Harrogate is about N.E. and not nearly E., as this would lead us to infer. Its course across the moors from Bolton Abbey to Harrogate is well marked on the high ground about Beamsley Beacon, but thence eastwards it is not so apparent until the lower beds are brought up within a mile or two of Harrogate, when its general

outline again becomes very distinct.

At Harrogate the general geological structure is rendered very clear by the massive grits at Pannal and Birk Crag, where the same bed of grit, striking in a north-east direction, forms, as it were, a wall on either side of the anticline, and gives at a glance the key to the main geological features of the neighbourhood. The axis of this disturbance is well shown in the town of Low Harrogate by the quarries that have been opened for roadstone on either side of the Bogs Field, where the rock may be seen dipping respectively N.W. and S. E. at high angles on either side of this valley; the same thing occurs near Beckwith Shaw to the S.W. of Harrogate. From these two positions the general run of the axis of disturbance may be easily gathered.

The exact base of the Millstone Grit is not so easily traceable in all cases, and can only be followed by very close inspection and consideration of the country. On the southern side of the anticline its general course is from the south of Pannal Ash, across the Stray to High Harrogate; to the north of the town its outcrop being obscured by faults is not so well marked, the junction with the underlying measures ranging along a N.E. line from near Shaw Green, to the north of Harlow Hill and Low Harrogate. Between these two lines is an area of about three square miles of lower measures upon which the town of Low Harrogate is situated.

The highest point of this area is Harlow Hill, 600 feet above the sea, and it is here that the greatest amount of disturbance appears to have taken place, The sandstone which caps this hill is one of the lowest beds of the district; and underlies the roadstone of Harrogate and Beckwith Shaw, towards both of which places it has a considerable dip. About a mile on either side of this hill occur most of the principal Sulphur Springs of the district.

Let us now turn to the chemical side of the question, and see what are the principal constituents of the waters. A very large number of analyses of these waters have been published from time to time, as may be seen from the list of works given in the bibliography. Those given on the opposite page are the most recent that we have been able to come across. These show that the principal constituents contained in these waters enable them to be separated into two distinct groups, the one containing an alkaline sulphide and the other a proto-salt of iron. These may be divided into a series of sub-groups or classes according to their chemical constitution, and the amount of saline matter they contain.

#### GROUP I.—SULPHUR WATERS.

	(	Strong	{	The Old Sulphur Water. The Strong Montpellier Sulphur Water
Saline Sulphur -		Mild	$\bigg\{$	The Mild Sulphur Water. The Mild Montpellier Sulphur Water. The No. 36 Water. The Magnesia Water. The Crescent Saline Water.
Alkaline Sulphur	-	the .	$\left\{ \right.$	The Starbeck Sulphur Water. The Beckwith Sulphur Water. The Harlow Car Sulphur Water.

#### GROUP II.—IRON WATERS.

Saline Iron	ee .	The Kissingen Water. The Chloride of Iron Water. The Alum Well. The Alexandra Water.
Pure Chalybeate-	-	The Tewit Well. The Johns Well. The Pure Chalybeate Well

No. 36. Sulphur Well.	26.4 26.4 26.4 241.6 11.1 7.2 11.1 7.2 14.529 1.30.4 26.6 1.30.5
Beckwith Spring.	6-10 trace 6-10 trace 3:374 
Starbeck Spa (Fairley).	1.515 trace  1.0070 trace 109.890 trace 109.890 trace 0.001 0.225 7.825 7.825 7.825 7.825 1.745
Magnesia Well (Muspratt),	1.222 trave trace trace trace trace trace trace trace trace trace trace 12.799 12.799 12.799 11.50
Mild Sulphur Well Montpellier (Attfield).	888.800 97.589 5.691 656 388.800 656 388.800 656 388.800 656 388.800 656 388.800 656 838.800 656 838.800 656 838.800 656 838.800 656 656 656 656 656 656 656 6
New or Mild Sulphur Well Royal Pump Room (W. A. Miller).	6.89 trace 16.70 2.39 11.34 trace 582.95
Strong Sul- phur Well Montpellier (Attfield).	14.500 14.500 79.936 57.989 4.811 trace 996 827.371 - - - - - - - - - - - - -
Old Sulphur Well Royal Pump Room (Thorpe)	6.566 trace 43.635 48.281 9.592 7.753 1.031 893.670 2.283 1.13
Saline Constituents in Grains per Gallon,	Sodium Sulphydrate - Sodium Sulphydrate - Barum Chloride - Strontium " "

TABLE OF ANALYSES OF THE SALINE CHALYBEATE AND PURE CHALYBEATE GROUP.

Alum Well (Davis).	69.33 778.76 89.47 56.91 57.33 3.14 2.19 83.96
John's Well, or Old Spa.	307
Chloride of Alexandra Carbonate of Pure Chaly. Iron Spa (Muspratt). (Muspratt). (Muspratt).	
Pure Chalybeate. Royal Pump Room (Davis).	1.625 trace trace trace 1.364 1.364 1.364 1.103 2.262 8.00 2.258
Carbonate of Iron Spa (Muspratt).	7.625 
Alexandra Chalybeate (Davis).	17.04 17.04 17.04 17.04 18.98
Chloride of Iron Spa (Thorpe).	
Kissingen Spa (Attfield).	21.425 674.588 674.588 65.439 1.590 2.136 8.858 8.858 8.858 8.858 8.858 1.50 2.136 8.858 8.858 8.858 8.858 8.858
Saline Constituents in Grains per Gallon.	Ferrica Sulphate Ferrica Aluminium "Calcium "Potassium "Ferrous Chloride Fotassium "Ferrous Chloride Fotassium "Ferrous Calcium "Maganese "Ferrous Calcium "Galcium "

From the above tables it is seen that the Harrogate mineral waters consist of four or five distinct classes, each of which has a certain marked peculiarity in its chemical constitution. The Strong Sulphur is very rich in chlorides and sulphides with a large bulk of free gases. The Mild Sulphur has nearly the same constituents but in a less concentrated condition. The Saline Chalybeate is very rich in chlorides and carbonates but contains no sulphides. The Pure Chalybeate contains a much smaller proportion of saline ingredients than either of the preceding; in fact it more nearly resembles the ordinary kinds of drinking water with a rather larger percentage of the carbonates of

magnesia and iron.

It is not our purpose, however, to go into the chemical constitution of these waters, but rather to note what connection there may be between so great a variety of classes of water and the geological structure of the neighbourhood. All the writers on this subject agree in considering that the mineral waters of Harrogate are owing to the peculiar geological structure of the district, but they differ somewhat in their ideas as to their origin. Some, and they are rather a majority, consider that most if not all the Harrogate mineral waters have a common source; and that the differences in the springs, which are only partial, are occasioned by the different channels through which they reach the surface. Others again consider that each spring has its own independent source, and that it exerts no influence over any of the neighbouring springs as long as the water remains below the surface. Professor Phillips in his description of the geology of this part of Yorkshire considered that the Chalybeate and Sulphur springs of Harrogate originate from the junction of two lines of fault, one running nearly north-east, and the other about due north\*. With regard to the first of these two lines of fault, it is very apparent that there is at least one if not several faults running in a more or less north-east direction and parallel to the main axis of disturbance; but from the manner in which the Harrogate anticline is enclosed both on the north and south sides by great ridges of grit striking across the country in an unbroken line for some considerable distance on either side of Harrogate, it is not possible that the strata can be broken by a north and south line without the dislocation being more apparent, and even supposing there were a cross dislocation there is no reason to suppose that it would influence the underground water at Harrogate any more than at hundreds of other places where the same thing occurs. It seems much more probable that the real source of these springs is from the same or nearly the same beds of shale as those from which they issue at Harrogate; that they are confined to this group of strata throughout their entire course; and that the faults and disturbances at Harrogate exert only a minor influence on the peculiarity of the waters. In the country west of Harrogate the strata in a general way consist of massive grits separated from one another by varying thicknesses

<sup>\* &</sup>quot;Illustrations of the Geology of Yorkshire," Part II., p. 253.

of shale which are thrown into a series of undulations, but are not broken by any faults of much importance. These beds, according as they partake of the sandstone or shaly character, become pervious or impervious to the passage of water, the sandstones allowing the percolation of water through their substance in all directions, whereas the shales admit its passage only in the direction of the planes of lamination, and only then when they have become tolerably arenaceous. From this it follows that water which issues from a certain orifice, say the Old Sulphur Well at Harrogate, would be confined throughout its entire course to the same strata, although probably not to exactly the same beds as those between which it springs at Harrogate unless the contiguity of the shales were broken by large faults, and these as we have noticed do not occur in the region to the west.

There is another circumstance to be taken into consideration, and that is the thinning out of some portion of the strata. The effect of this would be that, in the case of thinning out occurring towards the west, a stratum of water entering the ground in that region would become divided by the wedge-like nature of the rocks, so that two or more springs might still have the same source in the first case, but have become gradually separated in their passage eastwards. Supposing, on the other hand, that the thinning out took place in the opposite direction, the effect of course would be exactly the reverse; so that the water which entered the ground by several channels would become united into one stream by the gradual dying out of the intermediate rocks.

One great proof of the separation of the several springs at Harrogate is that the water in the wells, often only a few feet apart, stands at totally different levels, and that pumping from one well does not affect another; whereas if there were a common source for the water in the immediate neighbourhood, the same pressure would be exerted in all cases, the water would stand at the same level in all the wells, and the pumping the water in one well would in an equal degree affect all the neighbouring wells. If this be true, then, that each spring is confined throughout its entire course to the same, or nearly the same, geological horizon as that from which it issues, its origin will be where the same strata comes to the surface in the high ground to the west. This is, as we have stated, along the valley of the Wharfe, near Bolton and Appletreewick, and beyond, being about twenty miles or more due west of Harrogate. The passage of the water will be after the manner of an inverted syphon; pressure will be exerted by the superior elevation of the strata to the west, which will force water up through the highly inclined rocks at Harrogate.

Again, the temperature of the water forms what may be considered almost a direct proof that the sources of the wells are not superficial while at the same time it also proves that they do not spring from an enormous depth. Dr. Bennett observes "that the temperature of these sulphur springs during the

summer was always above that of springs of fresh water in the vicinity, and that in winter, as in the month of January, when the thermometer has fallen to 29 degrees, these springs were never lower than 41 degrees."\*

This shows that the springs are deeply seated enough to be somewhat affected by the internal temperature of the earth, but not so deep as to be beyond the influence of the external temperature of the air, or so deep as to become what are termed warm or hot springs, like those of Bath and elsewhere. Of course this does not apply to the simple Chalybeate waters, which have no peculiar features other than those common to the ordinary class of springs, and some of which are freely acted upon by changes of temperature.

The next point to consider is the different qualities of the water, and how this difference is occasioned. For this purpose the division into classes is exceedingly useful. The geological map shows that all the Sulphur waters with the exception of those at Starbeck issue from some of the lowest strata that occur in the district, and that the Strong Sulphur wells are those nearest to the axis of the anticline, while the Mild Sulphur and the Saline Chalybeates are at a slight distance on either side. The Pure Chalybeates rise irrespectively from both high and low measures, although some of the principal springs of this water issue from the highest strata, that is the Millstone Grit.

Each class of springs, then, in a general way issue from different groups of strata, but there are apparently some exceptions to this rule, arising in many cases, no doubt, from our ignorance of the original orifice of the spring. This may be at some distance from where the water now issues, the latter part of its course having been diverted by superficial material as well as by faults and thrusts in the underlying strata. The effect of the mixing of two classes of water is shown in the springs about the Bogs Field, where the sulphuretted hydrogen of the sulphur water combines with the iron from the Chalybeate springs, and produces the black, dirty-looking water of these wells. Accepting the hypothesis, then, that each class of springs issue from, and are confined to, distinct bands of strata throughout the greater part of their entire course, it does not seem difficult to account for the difference in quality of each of these groups; but if, on the other hand, we assume that there is a common repository for all the wells, and that the water reaches the surface by faults and fissures in the strata, why is it that there is so much difference between these various classes of water? for surely if the supply were drawn from all sorts of strata indiscriminately, the chemical peculiarities would become mingled, and there would be nearly the same constituents in all the springs. Again, if the chemical impregnation of the water took place in its passage to the surface through joints and fissures, the sources from which it would derive its supply of salts must be comparatively near

<sup>\* &</sup>quot;Observations on the Sulphurous Springs of Harrogate," 1843, p. 1.

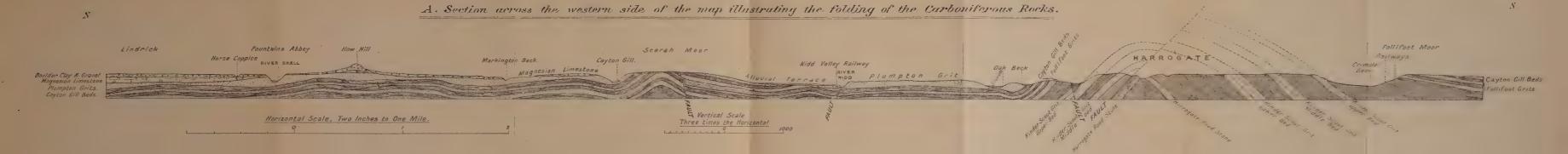
the surface, the amount would be but limited, and the quantity dissolved by the water would perceptibly diminish in course of time. But this is not the case, for several of these wells are known to give off an enormous amount of saline matter, which was not decreased to any extent since analyses were first made; Mr. Thackwray's wells alone producing as much as ifteen tons

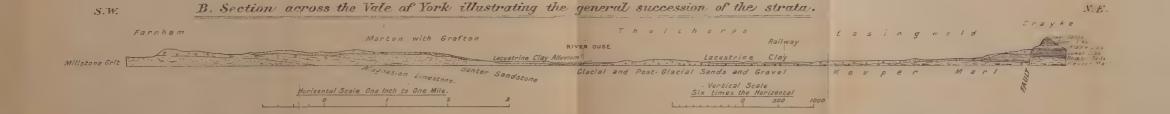
of salt in the year.

The stores from which these springs derive their saline matter must extend over a large area, and such an area is afforded by the Lower Carboniferous rocks, which occur at a short distance below the surface west of Harrogate. There are throughout this area undoubtedly several instances of the thinning out and alteration of the different beds, but they probably occur at some little distance from Harrogate, and would in no way affect the distinct qualities of the water, which would not have been acquired until they had become confined to distinct channels. The chemical peculiarities of the water are such that there is no reason for supposing any source other than what these shales could supply when exposed to the action of Mr. Hudleston considers that the deoxidation of the sulphides is caused by the decomposition of vegetable matter. "The considerable quantity of carburetted hydrogen, and of nitrogen . . . . favours this view, which is much strengthened by the almost complete absence of oxygen, showing that the nitrogen present is due to organic decomposition."\*

Taking all these things into consideration, it seems the simplest and most probable theory, for theory it must be; firstly, that each spring has its own independent source, with which it is connected by a separate and distinct channel formed by the alternating series of pervious and impervious strata which connects Harrogate with the hilly regions to the west, for it is only from such a source as this that a large body of water could flow year after year with such slight alteration both in quality and quantity; secondly, that the chemical impregnation of the water is caused during its passage through these strata, and that it is from this source that the bases of all the salts are derived.

<sup>\*</sup> Proc. Yorksh. Geol. Soc., vol. viii., p. 123.







### APPENDIX I.

# BORINGS AND WELL SECTIONS.

# Harrogate.

Well at Knox Farm.
(One-inch Map 93, N.W., N.S. 62; six-inch Map 154, S.W.)
About 325 feet above Ordnance Datum.
Communicated by Mr. Haigh.

Coarse yellow sandstone Blue shale with thin beds of stone	Thickness. Ft. in 75 0 - 90 0	Depth. Ft. in. 75 0 165 0										
Boring at the Ripon and Skipton About 225 feet above Ordnance Communicated by Mr. F. J.	Datum.											
Soil	- 6 0	$\begin{array}{ccc} 1 & 6 \\ 7 & 6 \\ 32 & 6 \end{array}$										
Boring on road 200 yards N. of crossroads. About 260 feet above Ordnance Datum.												
Yellow sand		$\begin{array}{cc} 10 & 0 \\ 60 & 0 \end{array}$										
Boring on road 380 yards N. of e About 250 feet above Ordnance	Datum.											
Soil	- 8 0	2 6 10 6 55 6										
Boring about 400 yards N.E. of c About 325 feet above Ordnance	Datum.											
Yellow sand Boulder clay	- 10 6 - 3 0 - 40 0	10 6 13 6 53 6										
Boring on road about 230 yards E. o About 300 feet above Ordnance	f crossroads. Datum.											
Soil - Yellow clay	- 3 0 - 2 6 - 10 6	3 0 5 6 16 0										
Blue shale		19 0 20 0 25 0										
HIGH HARROGATE												
Well between the Dragon Hotel and Communicated by Mr. H.	Lee.											
Hard gritty rock	25 0	58 0 67 0 92 0 102 2										

#### Starbeck.

(One-inchi Map 93, N.W., N.S. 62; six-inch Map 154, S.E.)

Boring at S. end of Bogs Lane. About 320 feet above Ordnance Datum.

									Thickness.	Depth.
									Ft. in.	Ft. in.
Clay			**		<u>.</u> .				11 0	11 0
Red sand and clay	mix	ed	-	-	-	-	-	-	9 7	20 7
Red ferruginous s				_	-	-	-		3 0	23 7
Strong red earth					-	-	-		14 6	38 1
Strong blue bind	-	-	-	-	-	-			2 6	40 7
Soft stone minera		-		-		_		-	11 4	51 11
Blue bind -	-	-	-	-	-	-	-	-	7 8	59 7
Black shale -	-				-	-	-	-	5 7	65 2
Galliard	-	-	-				-	-	1 0	66 2
Blue bind -	~	-	-	-			-	-	105   0	171 2
Galliard	-	-	-	-		-	-	-	0 3	171 5
Strong blue bind	-	-	-	-	-	-	-		€ 8	178 1
Black shale '-		-	-	-	-	-	-	-	3 3	181 4
Stone bind -	-	-	-	-	-	-	-		0 11	182 3
Black shale -	-	-	-	-	-	-	-	-	117 0	299 3
Stone bind -	-	-	-	-	-	-	-		0 9	300 0
Strong open black	shal	le	-		-	-	-	-	$5   8\frac{1}{2}$	$305 8\frac{1}{2}$
Black shale with									131 5	$437  1\frac{1}{2}$

Boring on road 100 yards W. of railway crossing. About 275 feet above Ordnance Datum. Communicated by Mr. F. J. Dixon.

					~					Thick	ness.	Dep	oth.
										Ft	. in.	Ft.	in.
Alluvium		-	~	-	×	~	-	-	-	5	0	5	0
Sand-		-			-	-		-	-	7	0	12	0
Running s	and -		-	-	-	-	-	-	-	10	θ	22	0
Grey sand	lstone	-	-	-	-		-	-		6	0	28	0
	Boring	near Abo	r foo	tpath 75 fee	1 330 t abo	yard ove C	s S. Irdna	of ra	ilway Datu	r crossi m.	ng.		
Alluvium	[Drift?]	-	-	-	_	_			_	4	0	4	0
Sand		-		-		-		-		6	0	10	0
Soft sands	tone		-	-	-	-		-		16	0	26	0

# Spofforth.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 171, N.E.)
Boring on Follifoot Ridge, about 1,000 yards S.W. of centre of tunnel.
About 425 feet above Ordnance Datum.
Communicated by Mr. F. J. Dixon.

									Thickness.	Depth.
613									Ft. in.	Ft. in.
Clay									7 0 .	7 0
Blue shale -		-			-	-	-	-	33 0	40 0
Rock and shale		-	~	-	-	-		-	47 0	87 0
Hard sandstone				-		-	_	-	4 0	91 0
Sandstone -		-		10	-	_	-		19 0	110 0
Black shaly rock	-									114 0
Sandstone grit -		-			_			_	17 0 -	131 0
Hard rock -					-				3 0	134 0

#### KIRK DEIGHTON.

Well in shelly limestone, about 30 yards.

### IGMANTHORPE HALL.

Well about 21 yards to limestone. About 70 or 80 yards from this got water at 7 yards, but met with no limestone.

### Bilton.

THE HALL.

(One-inch Map 93, S.W., N.S. 70; six-inch Map 173.) Sand about 20 yards, then Sandstone.

Well between Jewitt House and Blacksmith Arms opposite end of lane to Healaugh. Communicated by Mr. Montague.

								Thickness.	Depth.
								Ft. in.	Ft. in.
Fine soil .		-	-	-				- 1 0	1 0
Rubble	-	-				_		- 3 0	4 0
Strong red sand	*		- ,	_	-	-	about	59 0	63 0
Red sandstone									

Good supply of water obtained.

#### WIGHILL.

Wells about 8 or 9 yards deep. Water from gravel.

#### Burton Leonard.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 137, S.E.) Boring 760 yards W. of centre of village. About 300 feet above Ordnance Datum. Communicated by Mr. F. J. Dixon.

								Thickness. Ft. in.	Depth. Ft. in.
Boulder clay	-	-	-	-	-	- '		12 0	12 0
Marl and beds of rock	-	-	-	-	- '	-	-	28 0	40 0
Hard limestone -	_	~	-	-		-		184 0	224   0

# Staveley.

Well at Loftus Hill. (One-inch Map 93, N.W., N.S. 62; six-inch Map 138, S.W.)

(0.20 12		r		.,		,			Thiel	eness.	Dept Ft. i	
Red earthy clay		_	_		-	_		_	12	0	12	
Broken limestone	gettin	ng st	ronge	r tov	vards	botto	m	-	66	0	78	()
Red grit like that	of K	nares	shorou	ıgh								

### ALLERTON HALL.

Gravel, 25 yards.

#### ARKENDALE.

Clay, 30 feet, then gravel.

#### Brearton.

Well in the village.

			23 1)	out 1	10 10	560	cus	0016	Otune		A) EUD 14.	Thiel	cness.		
												Ft.	in.	Ft. i	n.
Clay	-	ata .	_		-		_	-	-	**	-	9	0	9	0
Fravel										-	-	1	2	10	2

## Ingerthorpe.

Well at Mr. Ward's House, Jan. 1864.

Comn	nunic	ated	b <b>y</b>	Rev.	J. S. Tute	Thickness. Ft. in.	Depth. Ft. in.
Brown gravelly clay -		-	-	- 4	about	6. 0 .	6 0
Dark brown boulder clay Blue Magnesian Limestone				-	about	39 0	45 0

# Markington.

Well in the square. Communicated by Rev. J. S. Tute.	Thickness.	Depth. Ft. in.
Stiff dark-coloured boulder clay	21 0	21 0
Sandy gravel with boulders included rounded pieces of Magnesian Limestone 9 0 t	o 12 0	33 0

### Ripon.

1. Well at Mr. S. Darnborough's, Hutton Bank. Communicated by Rev. J. S. Tute.

Thickness.

Depth. Ft. in.

										T. f.	LLI .	T 00 1	LLL
Soil				-	_	-	-	_	-	0	8	0	8
Soft sandy red	roc	k -		-	-	4	-	-	-	10	0	10	8
Soft marly cla		-	-	-	-	-	**		-	0	10	11	6
Red rock mi	ich	false-	bedd	ed v	with	laye	rs	of c	elay				
about eve	ry 10	) feet		-	-	-	-		-	74	0	85	6
			2	2. Bo	ndga	te Bı	ewe	ery.					
										Thick	ness.	Dept	th.
										Ft.	in.	Ft. i	in.
Sand and grav	rel	-	-	-	-	-	-	-	•	30	0	30	0
Limestone	-	-	-	-		-	-	-	-	2	0	32	0

BLOWS HALL. Red sand, 47 yards.

MARTON-LE-MOOR. Sandstone at 70 feet.

HIGH BROOMS.
Sandstone at 75 feet.

NEWBY HALL. Clay, 7 yards, then wick sand.

# Boroughbridge.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 138.)
Well in centre of St. James Square.
From Brit. Assoc. Rep. for 1876, p. 107.

Thickness. Ft. in.	Depth. Ft. in.
Soft red sand and boulders 28 0	28 0
New red sandstone with layers of marl 3 in. to 4 in.	
thick 228 0	256 0
Water stands at 17 feet. Reduced 2 in. after 36 hours pumpin	r.

### Brafferton.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 120.)

1. Boring at Helperby.

Communicated by Prof. Kendall.

										Thickness.	Depth	1.
										Ft. in.	Ft. ir	1.
Boulder clay	and	sand	-	-	-	-	-	-		20 0	20	0
Quicksand	**	**		**	-	-				10 0		ŏ
Fine clay			~	-		_	_	_		10 0		0
Gravel -						_	_	_		15 0		0
Sandstone			٠,	_						19 0	99	U
Danustone			-			-		-	-	-	- April	

# Brafferton—(continued).

				2	. В	oring	at P	ill Mo	or.		Thick	ness.	Depth.	
~				Comm	uni	cated '	by A	Ir. <b>T</b> .	Ow	ton.			Ft. in.	
Soil and sa	and	-		-	_	-	~	-	-	~	3	0	3 0	
Clay	~		-	-	-	-	-	-	-	-	115	0	118 0	

# Sessay.

Sand, 7 yards.

# Myton.

Sand, 10 yards.

## Raskelf.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 121.)												
1. Spring House.  Communicated by Mr. T. Owston.												
Boulder clay 60 0 Freestone rock 145 0  Water rises within 6 feet of the top.	s. Depth. Ft. in. 60 0 205 0											
2. Boring near the railway station.												
From Rep. Yorksh. Phil. Soc. for 1893, p. 57.												
Sand	10 10 50 10 65 10 106 10 151 10 161 10											
Huby.												
(One-inch Maps, 93, N.W., N.E., N.S. 62, 63; six-inch Maps 1  From Rep. Yorksh. Phil. Soc. for 1893, p. 57.  Thicknes Ft. in.												
Soil	0 6 13 6 25 6 105 6 109 6											
2. Huby Burn Farm.												
From Rep. Yorksh. Phil. Soc. for 1893, p. 56.												
Soil	1 6 11 6 19 6 23 6 35 6											
Green marl with gypsum Brown marl with gypsum Red sandy marl												

### Newton-on-Ouse.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 156.) Communicated by Mr. J. Villiers. Also in Kaye's Report. Well sunk, March, 1888.

				44 611	Sum	9 141 00	Thicks	ness.	Dept	Depth.			
										Ft.	in.	Ft.	in.
Top soil -		-	-	_			-		-	1	0	1	0
Red clay -	-		-		-	-	-		-	3	0	4	0
Warp -	-	-		-	-	-	-	-		7	0	11	0
Sand -	-	_	-			-	_	_		25	0	36	0
Sand and war	p	- ~		_		_	-	-		20	0	56	0
Service 1 1		-				÷ .	_	-		16	0	72	0
37 11 7		-	_		_	-	-	-	_	1	0	73	0
Light sandsto	ne		_			**	_	-	-	17	0	90	0
Soft red sands		e -		and the same		_	-		_	24	6	114	6
Red marl -	_	_					_	_	_	0	6	115	0
Soft sandstone	e	_		_		_	_	_		105	0	220	0
Water lev		32 fee	t fr	om sur	face								

#### Cattal.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 172.) At the Inebriates' Home, 1903. From Kaye's Report.

										Thickness. Ft. in.	Depth. Ft. in.
Soil	er	-	-	-	-	-	-		-	1 0	1 0
Boulder clay	-	-	-	-	-	-			-	17 0	18 0
Sandstone	-		-	-	-	-		-	-	184 0	202 0

### Kirk Hammerton.

(One-inch Map 93, N.W., N.S. 62; six-inch Map 156.) Well at the Railway Station.

							Thickness. Ft. in.	Depth. Ft. in.
Superficial beds	-	~	-		-	-	24 0	24 0
Red sandstone	-	-		-	~	-	er-tune	

### Thirsk

(One-inch Map 96, S.W., N.S. 52; Six-inch Map 87, S.E.)

Boring at the Brewery, 1877.

Communicated by Mr. J. T. Hymas.

						_		,	,	Thickness.	Dep	th.
										Ft. in.	Ft.	
Sand -		-		-	-	-	-	-		6 0	6	0
Clay -	-		-	-	-		-		-	4 3	10	3
Gravel with	water	***	-	-	-			-	-	6 9	17	0
Stony clay	-	-	-		-	-	**			5 10	22	10
Sand and gra	vel			-		_	÷.		~	4 3	27	ĭ
Brown strong	g clay	-	-		-	- `	-	-	-	3 1	30	2
Red shale	46.	-				-	-		**	1 6	31	_
Grey shale			-		-	-	-		-	18 10	50	6
Grey shale Soft red mar	l -	-	~	-	-	-	-	-	-	3 0	53	6
Red and grey	' marl	4		-	-	-			_	5 0	58	6
Grey shale an	d gir	dles	with	wate	r	-			-	5 6	64	0
Red and grev	shale	1	-		-		-	_		6 6	70	6
Shale and gyr Grey metal a	sum		-	**		-	-	-		10 6	81	0
Grey metal a	nd gy	psun	1	-		-	-	_		3 6	84	6
Red shale and	l gyps	sum	-		-		-	67	_	6 7	91	1
Grey shale a	ad gir	dles	with	gyps	sum	- 1				8 5	99	6
Red shale and	dgyps	sum			-	_		-	-	52 6	152	0
Grey shale as										10 0	162	ő
Red shale and										10 0	172	0
Grey shale an										$\tilde{3}$ $\tilde{2}$	175	2
Strong red sh	ale w	ith h	nard	girdle	38	_		_		18 0	193	2
9										20 0	199	Las

n n	hizal-	100		7 \			
	hirsk					Thickness.	Thomas.
Hard grey post Red shale and gypsum with Red shale and freestone gi Strong shale and freestone g Shale and gypsum with wat Grey shale and gypsum Red shale and grey post wi Soft red freestone Strong red freestone						Ft. in.	Depth. Ft. in.
Hard grey post		-		1		6 5	199 7
Red shale and gypsum with	water	-	-		-	29 4	228 11
Strong shale and freestone gir	rdles -	-	-	*	**	13 0	241 11
Shale and gynsum with was	ter -	-		-	_	$\begin{array}{ccc} 20 & 1 \\ 2 & 11 \end{array}$	262 0 264 11
Grey shale and gypsum -		-	_		_	15 0	279 11
Red shale and grey post wi	ith wate	91 -		-	-	11 0	290 11
Soft red freestone		-	-	-	-	10 9	301 8
Soft red freestone -		-	•	-	-	11 3 11 9	312 11 324 8
Red freestone		_	-		-	17 0	341 8
Red shale and freestone gi	rdles -	-		-	-	9 5 9 6	351 1
Soft red freestone - Strong red freestone - Red freestone - Red shale and freestone gi Red freestone -		-		-	-	9 6	360 7
	I	Beda!	le.				
	LOND	OND	ERRY	7.			
(One-inch Map 96					Mo	n 70 S W	
	unicated					,	
		3				Thickness.	Depth.
67 7 7 1						Ft. in.	Ft. in.
Clay and sand with water Red sandstone		-		~	-	36 0 60 0	36 0 96 0
Water is					ston		90 0
		LE E					
Sand and gravel with water		-			*	18 0	18 0
Sandy limestone		_	-	-	-	144 U 38 O	$ \begin{array}{ccc} 162 & 0 \\ 200 & 0 \end{array} $
Sand and gravel with water Boulder clay	elow su	rface.	Wat	er is	Iraw	n entirely fro	m
	the li	mestoi	ne bed	l.		J .	
	LEE	MING	BAI	₹			
	121717	MILLIO	10211			Thickness.	Depth.
	Boring a	t Mr.	Langs	taffs.		Ft. in.	Ft. in.
Sand and gravel with water	Boring a r	t Mr.	Langs	taffs.	-	Ft. in.	Ft. in.
Sand and gravel with water Boulder clay	Boring a	t Mr.	Langs	taffs.	-	Ft. in. 20 0 20 0 20 0	Ft. in. 20 0 40 0
Sand and gravel with wate Boulder clay - Red sandstone and red mar Red marl -	Boring a	t Mr.	Langs	taffs.	-	Ft. in. 20 0 20 0 20 0 83 0	Ft. in.
Sand and gravel with wate Boulder clay Red sandstone and red mar Red marl Water is	Boring a	t Mr.	Langs the re	taffs.	- - dsto	Ft. in. 20 0 20 0 20 0 83 0	Ft. in. 20 0 40 0 60 0
Sand and gravel with water Boulder clay - Red sandstone and red mar Red marl - Water is	sarawn	пош	one re	u san	- - dsto	Ft. in. 20 0 20 0 20 0 83 0	Ft. in. 20 0 40 0 60 0
v ater is	Mash	am	Moo	r.	usto	ne.	Ft. in. 20 0 40 0 60 0
(One-inch Map 9)	Mash Mash	am N.S. 8	<b>Moo</b> 51 : si	a san <b>T.</b> x-incl	asto 1 M:	ap 100 N.E.)	Ft. in. 20 0 40 0 60 0 143 0
v ater is	Mash , S.E., e Pott I	am N.S. 8	Moo 51; si Fair	er. x-inch Thor	asto 1 M:	ap 100 N.E.)	Ft. in. 20 0 40 0 60 0 143 0
(One-inch Map 9) Boring 400 yards S.E. of th	Mash , S.E., e Pott I	nom  N.S. & Beck at ance I	Moo 51 ; si 5 Fair Oatum	ct san T. x-incl Thor	n Ma	ap 100 N.E.) About 850 fe	Ft. in. 20 0 40 0 60 0 143 0
(One-inch Map 9) Boring 400 yards S.E. of th	Mash 7, S.E., e Pott F Ordn	nom  N.S. & Beck at ance I	Moo 51 ; si 5 Fair Oatum	ct san T. x-incl Thor	n Ma	ap 100 N.E.) About 850 fee Thickness.	Ft. in. 20 0 40 0 60 0 143 0 et above
(One-inch Map 9) Boring 400 yards S. E. of th	Mash 7, S.E., e Pott F Ordn	N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	ct san T. x-inch Thor J. Di	n Ma	ap 100 N.E.) About 850 fee Thickness.	Ft. in. 20 0 40 0 60 0 143 0 et above
(One-inch Map 9' Boring 400 yards S.E. of th  Comm	Mash 7, S.E., e Pott I Ordn unicated	N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	ct san T. x-inch Thor J. Di	n Ma	ap 100 N.E.) About 850 fee Thickness.	Ft. in. 20 0 40 0 60 0 143 0 et above
(One-inch Map 9' Boring 400 yards S. E. of th  Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Ft. in.  7 0 4 0 10 0	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0
(One-inch Map 9: Boring 400 yards S. E. of th  Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum fr. F.	ct san T. x-inch Thor J. Di	n Ma	Thickness.  Ttickness.  Tt. in.  7 0 4 0 10 0 12 0	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0 33 0
(One-inch Map 9: Boring 400 yards S. E. of th  Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Ttickness.  Th. in.  7 0  4 0  10 0  12 0  54 0	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0 33 0 87 0
(One-inch Map 9: Boring 400 yards S. E. of th  Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Ttickness.  Tt. in.  7 0 4 0 10 0 12 0	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0 33 0
(One-inch Map 9: Boring 400 yards S. E. of th  Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Thickness.  Ft. in.  7 0 4 0 10 0 12 0 54 0 8 0 0 6 3 0	Et. in. 20 0 40 0 60 0 143 0 et above E
Sandy soil  Brown shale  Brown shale  Brown shale  Brown shale  Brown shale  Hard blue grey rock  Brown shale  Hard blue grey rock  Brown shale	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Ttickness.  Tt. in.  7 0  4 0  10 0  12 0  54 0  8 0  0 6  3 0  4 0	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0 87 0 95 6 6 98 6 102 6
Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Thickness.  Ft. in.  7 0  4 0  10 0  12 0  54 0  8 0  0 6  3 0  4 0  10 0	Et. in. 20 0 40 0 60 0 143 0 1
Comm  Sandy soil  Brown shale  Brown shale  Brown shale  Brown shale  Brown shale  Brund grey rock  Brown shale  Blue shale (block)  Hard blue grey rock .  Brown shale  Hard brown rock	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Ttickness.  Tt. in.  7 0  4 0  10 0  12 0  54 0  8 0  0 6  3 0  4 0	Et. in. 20 0 40 0 60 0 143 0 et above E
Comm  Sandy soil  Brown shale  Hard grey rock Brown shale  Hard blue grey rock Brown shale  Yellow sandstone  Hard brown rock Hard brown shale  Yellow clay	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Thickness.  Ft. in.  7 0 4 0 10 0 12 0 54 0 8 0 0 6 3 0 4 0 10 0 3 0 10 0 0 9	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0 87 0 95 0 95 6 102 6 112 6 112 6 125 6 126 3
Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Thickness.  Ft. in.  7 0 4 0 10 0 12 0 54 0 8 0 0 6 3 0 4 0 10 0 3 0 10 0 0 9 5 9	Et. in. 20 0 40 0 60 0 143 0 1
Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  To use the control of th	Et. in. 20 0 40 0 0 60 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 144 0 144 0 145 6 126 6
Comm  Sandy soil Comm  Sandy soil Shade Shown shade Sh	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  Thickness.  Ft. in.  7 0 4 0 10 0 12 0 54 0 8 0 0 6 3 0 4 0 10 0 3 0 10 0 0 9 5 9 16 0 8 4	Et. in. 20 0 40 0 60 0 143 0 et above  Depth. Ft. in. 7 0 11 0 21 0 87 0 95 6 98 6 102 6 115 6 125 6 126 3 132 0 148 0 156 4
Comm  Sandy soil	Mash 7, S.E., e Pott I Ordn unicated	nam N.S. & Beck at ance I l by M	Moo 51; si Fair Patum r. F.	r. x-inch Thor J. Di	n Ma	Thickness.  To use the control of th	Et. in. 20 0 40 0 0 60 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 143 0 144 0 144 0 145 6 126 6

### Masham Moor—(continued).

#### ROUNDHILLS RESERVOIR.

Section in embankment trench, S. end.

Top water level 743. Stream level 630 feet above Ordnance Datum.

Communicated by Mr. F, J. Dixon,

					J.				Thic	kne	88	Der	oth.
									Ft.			Ft.	in.
Yellow sandstone		~	-	~	**	_	-	~	37	0			
Brown shale .	_		_		-		-	-	34	0			
Hard grey rock					-	-	~	-	12	0	to	18	0
Blue stratified sha	le	40	-		-		-	-	14	0			
Hard grey rock, st							-	-	2	0			
Blue stratified sha							-		22	0			
Thin bed of rock,									4	0			
Blue stratified sha													
out									20	0			
Blue stratified sh													
and thinning o									34	0			
								N. end	1.				
Yellow sandstone							_	,	13	0			
Brown shale -							_	_	35				
Hard grey rock									12				
Blue stratified sha									30				
Hard grey rock, u									90				
shale bands									79	0			
DIEUTO DUITUS			_						10	-			

# Kirkby Malzeard.

(One-inch Map 97, S.E., N.S. 51; six-inch Map 101 S.W.)
Boring on Grewelthorpe Moor. On the S. side of the stream flowing from Sandwith Wham and 430 yards S.W. of Stock Beck. About 875 feet above Ordnance Datum.

Communicated by Mr. F. J. Dixon.

									Thickness.	Depth.
~									Ft. in.	Ft. in.
Alluvium (Drift?)	-	-	-	-	-	-	-	-	49 0	49 0
Blue shale -	~	-	-	**	-	-	-	~	27 0	76 0
Grey sandstone			-		-		~	-	120 0	196 0
Blue shale -	-	-		-	-		-	-	21 0	217 0

Boring 930 yards S.E. of the above, and about 1,000 yards nearly W. of Bagwith House. About 800 feet above Ordnance Datum.

Communicated by Mr. F. J. Dixon.

										Thickness	Dept	h.
										Ft. in.	Ft. i	n.
Boulder clay	-		**	-	-	-		-	-	5 6	5	6
Blue shale		~	-	-	-	-		_	-	26 0	31	6
Grey sandston	ıe	-	~		-	-	_	_	_	61 0	92	6
Blue shale	-	-	-	-	-	-	-		_	9 0	101	6
Grey sandston	ıe.		_	_	_		_	-	_	12 0	113	6
Blue shale					**				_	30 0		-
THE SHOP							_		-	90 U	143	6

### APPENDIX II.

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List of Works relating to the Geology and Mineral Waters of the District with the exception of purely medical works.

Titles marked with an asterisk are taken from other bibliographies, but have not been verified.

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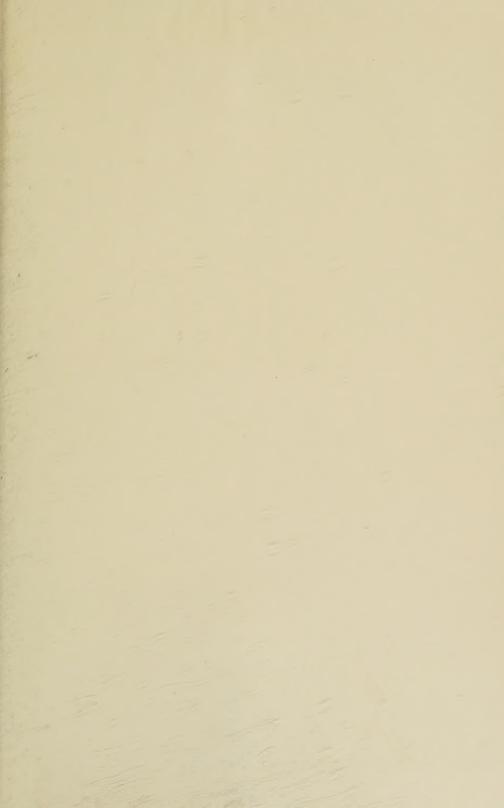
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